

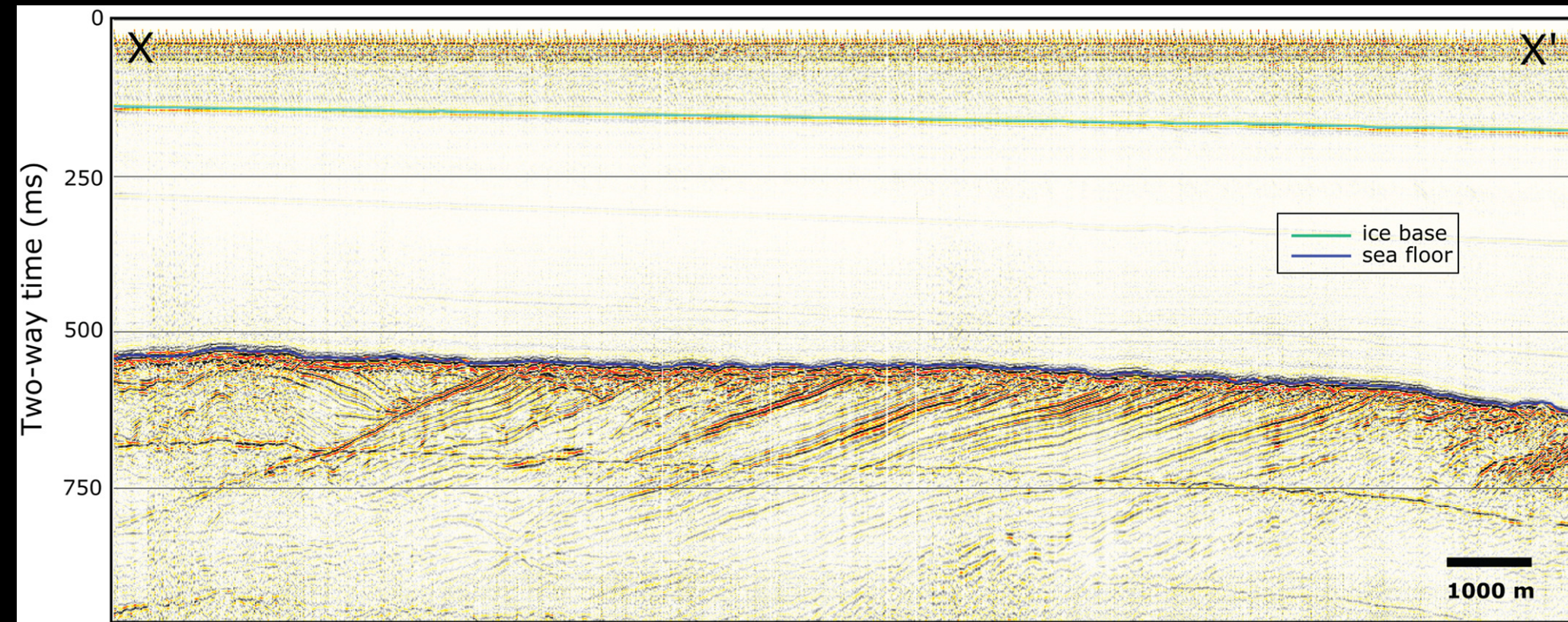




- Principles of Reflection Seismics
- Interval Velocities
- Basic (very basic) imaging principles



Smith et al., GRL, 2020

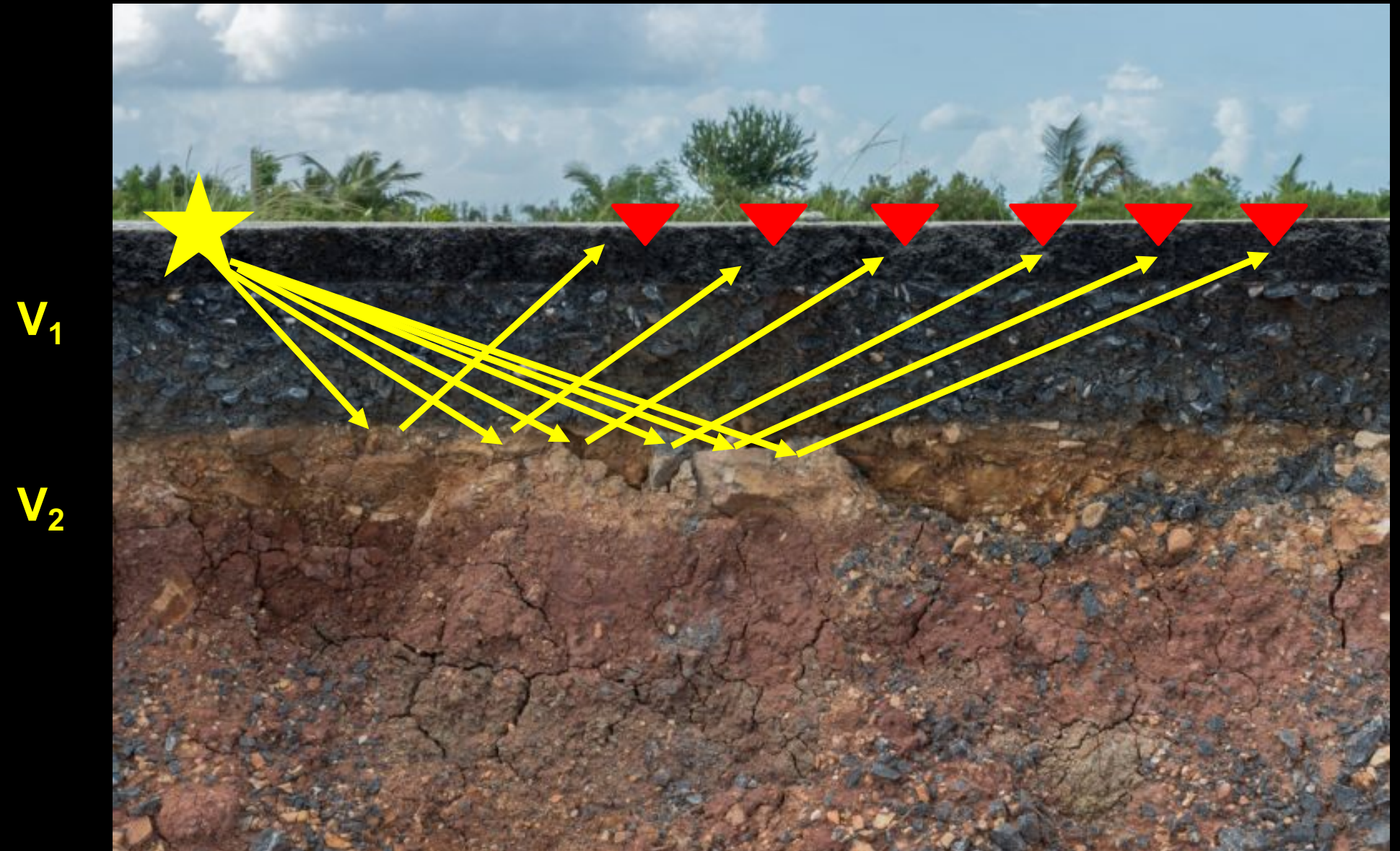


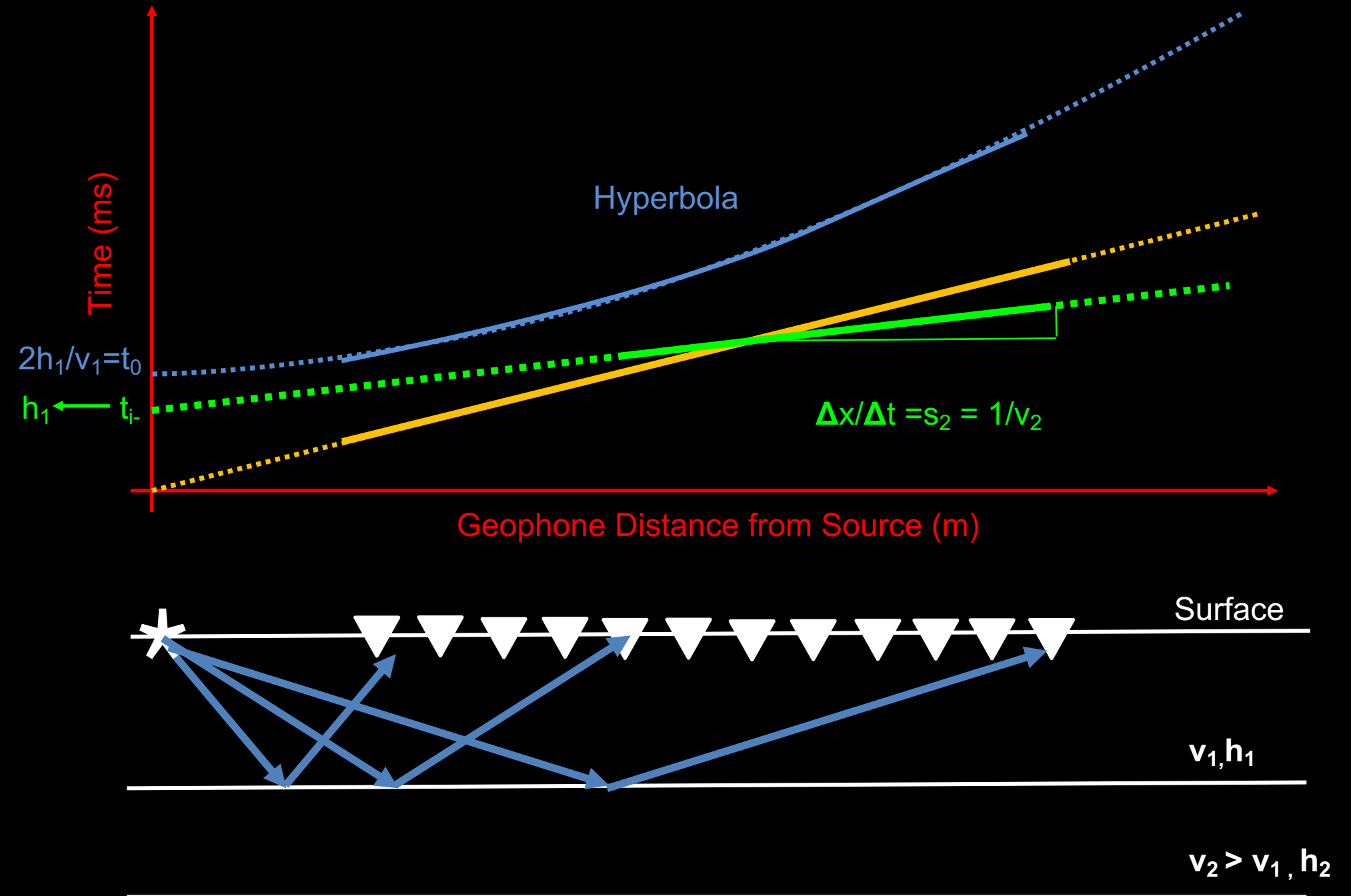


- Refraction seismics doesn't "image" the subsurface per se. It would be very hard, e.g., to find point targets etc.
- Reflection seismics can do this for you, but it takes a number of processing steps which can be delicate at times.
- We will try to get a peak into what is done.



Raypath for reflections

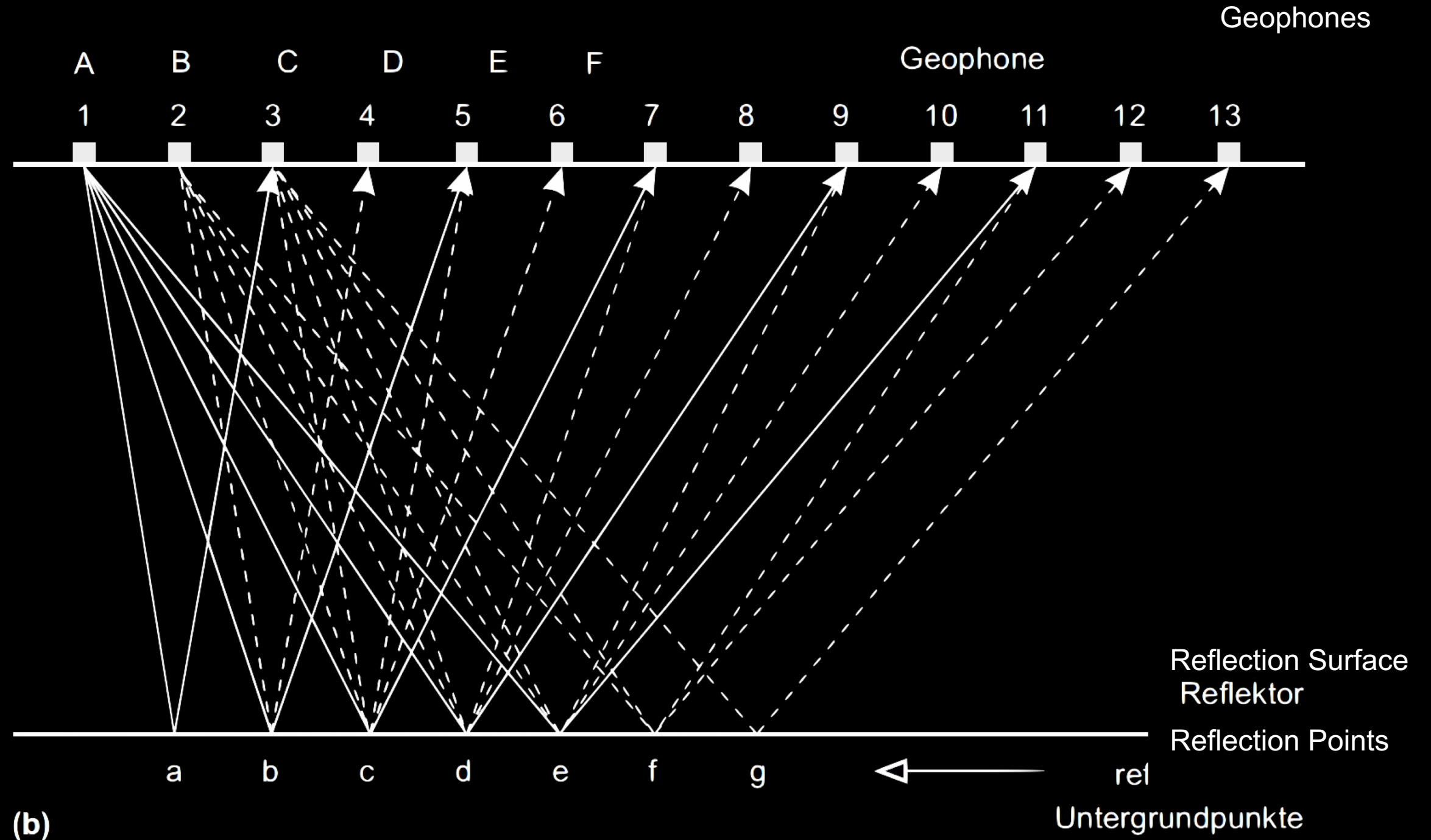






Reflection geometries for multiple shots

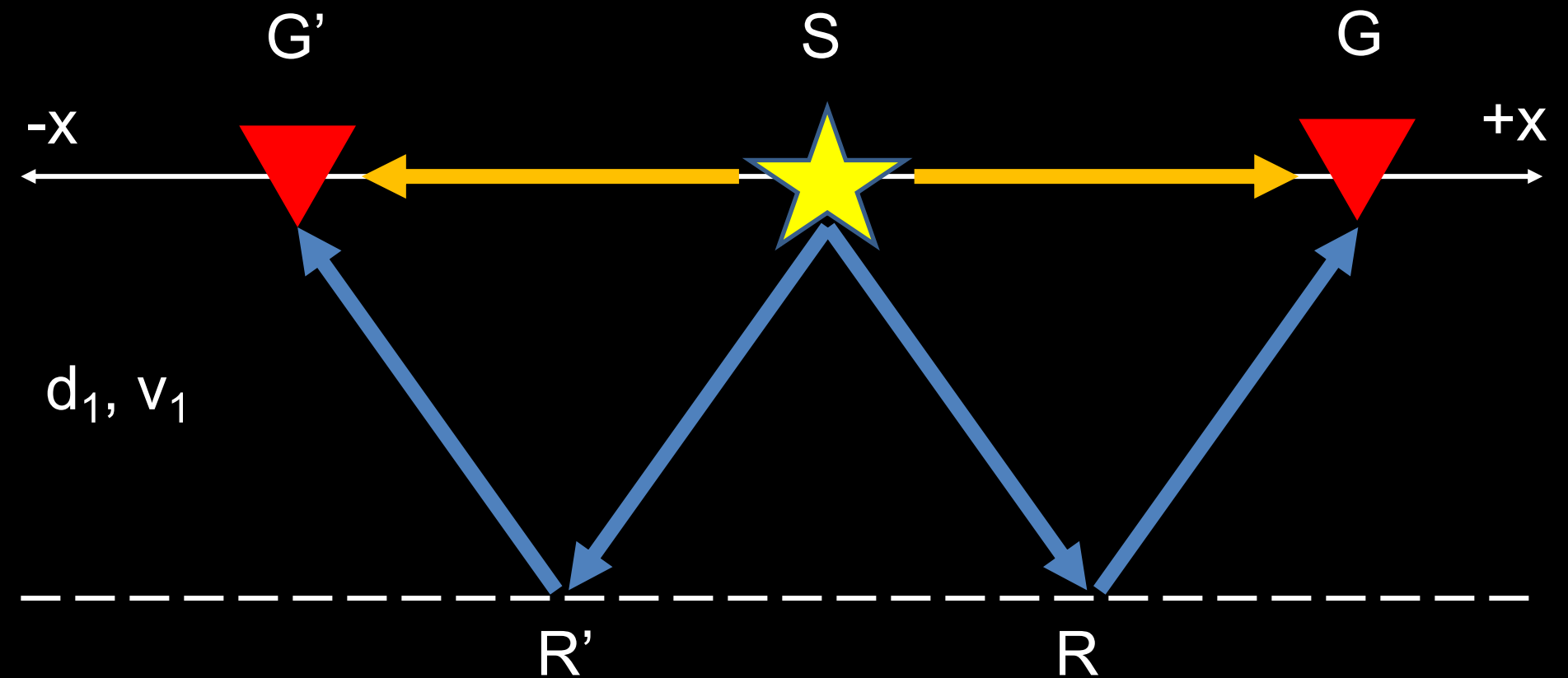
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- What are expected signatures in a single shot gather?
- What is the traveltimes of reflections?

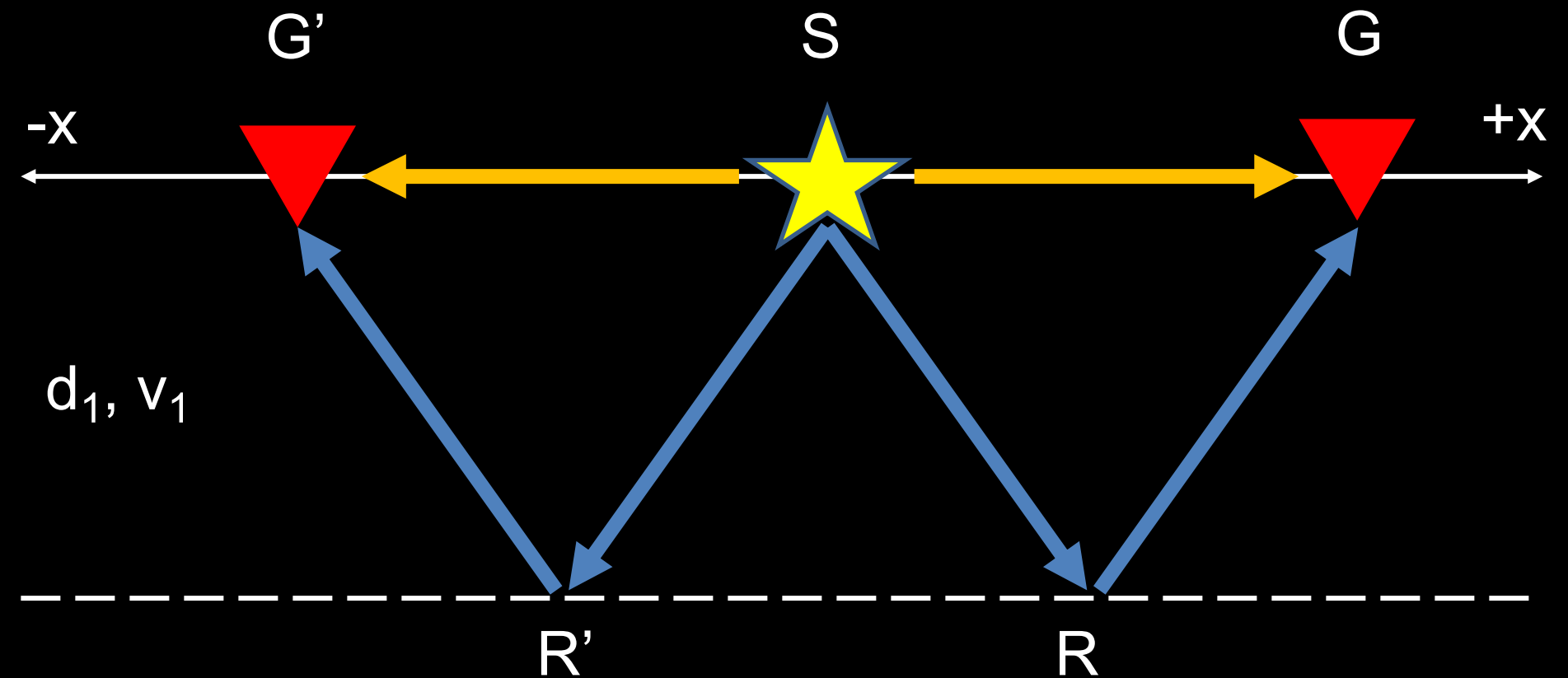


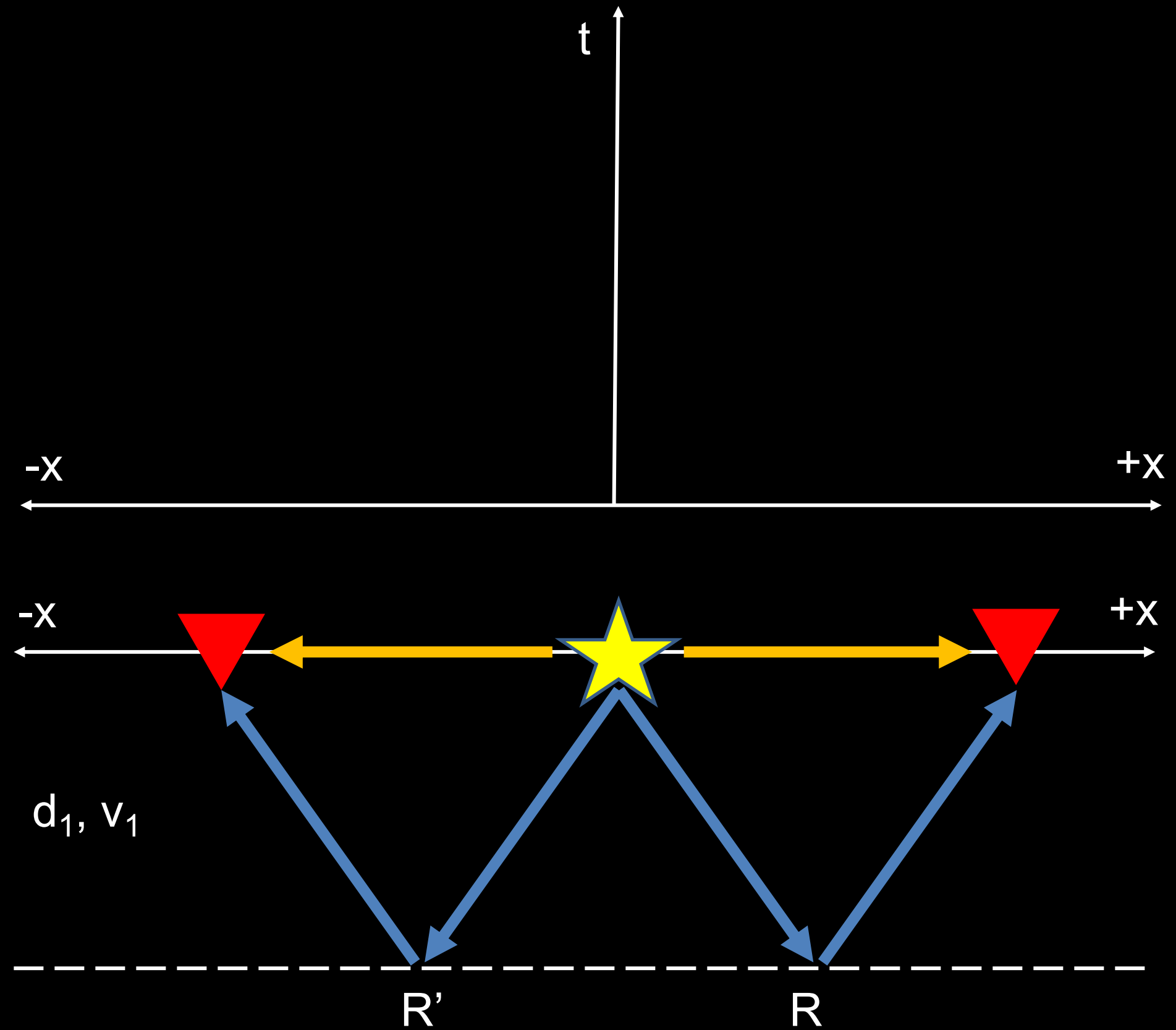


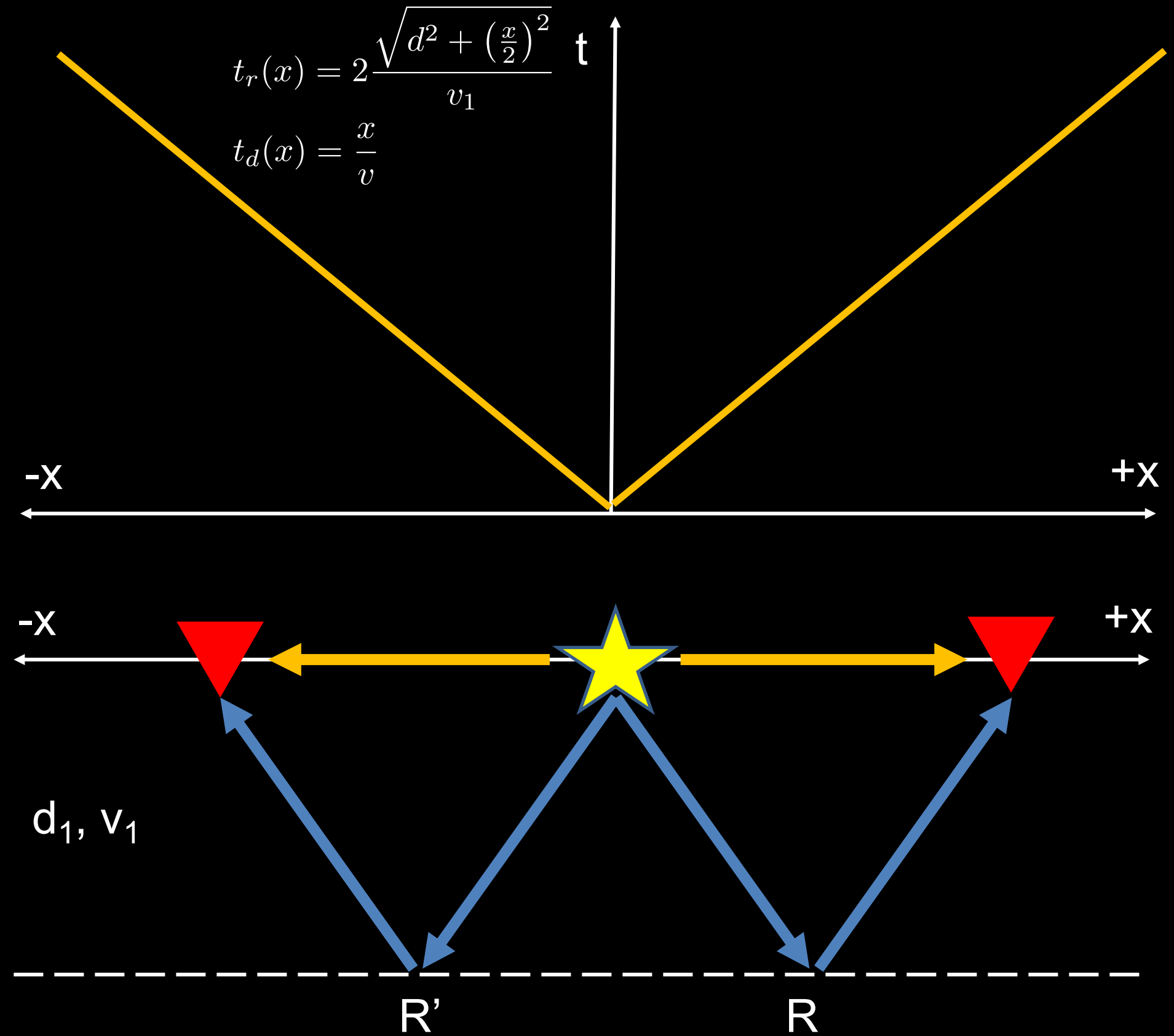


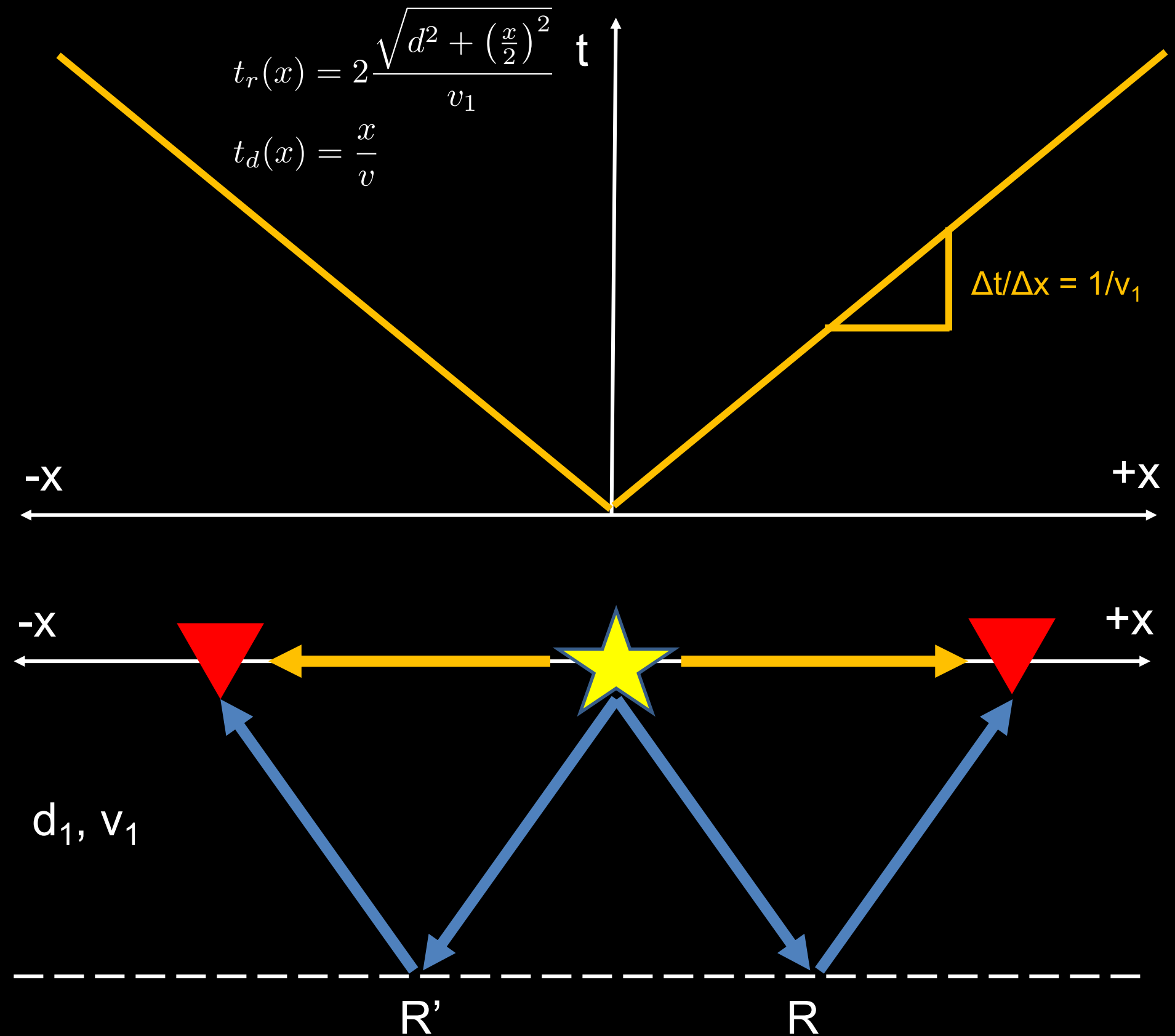
$$t_r(x) = 2 \frac{\sqrt{d^2 + \left(\frac{x}{2}\right)^2}}{v_1}$$

$$t_d(x) = \frac{x}{v}$$











Direct Waves

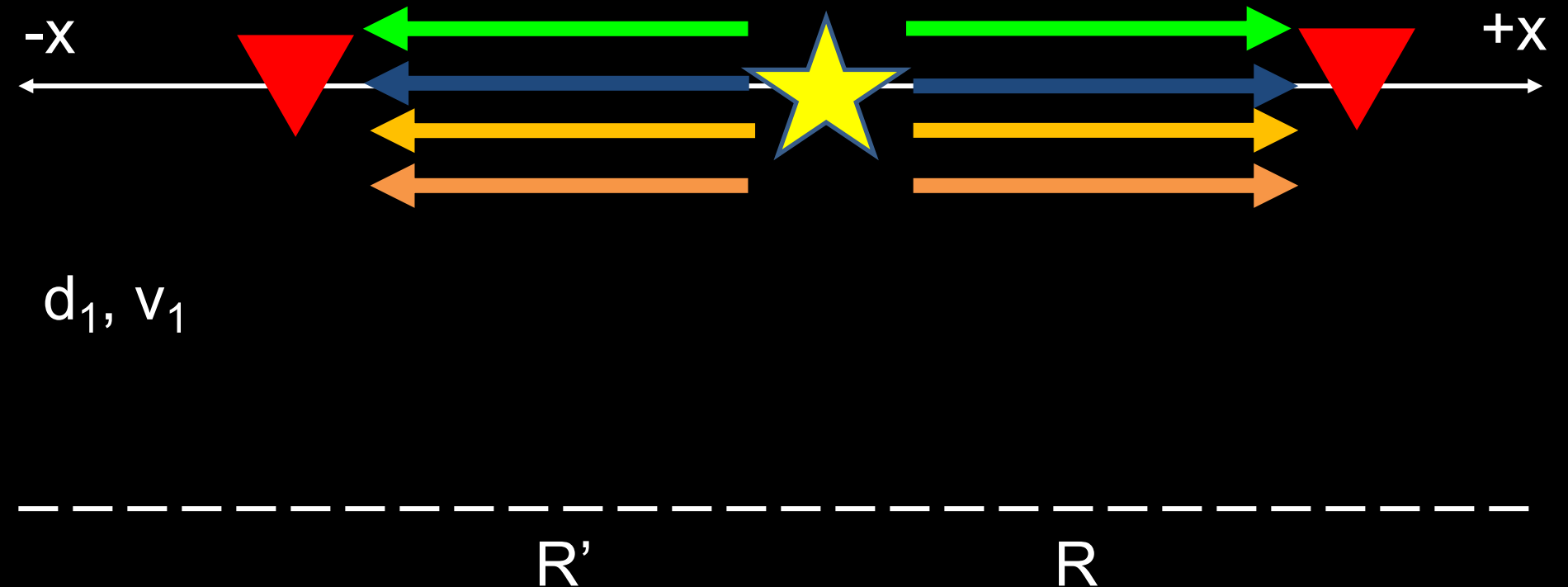
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Direct Air Wave

Direct Surface Wave (at boundary)

Direct s-wave (in body)

Direct p wave (in body)





Direct Waves

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Slow



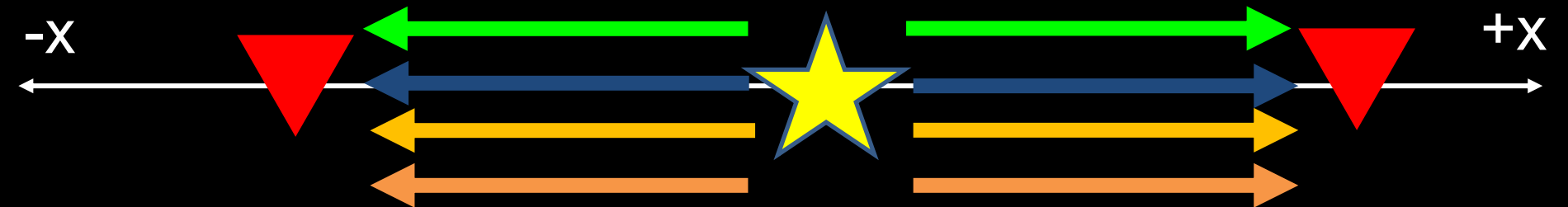
Fast

Direct Air Wave

Direct Surface Wave (at boundary)

Direct s-wave (in body)

Direct p wave (in body)



d_1, v_1

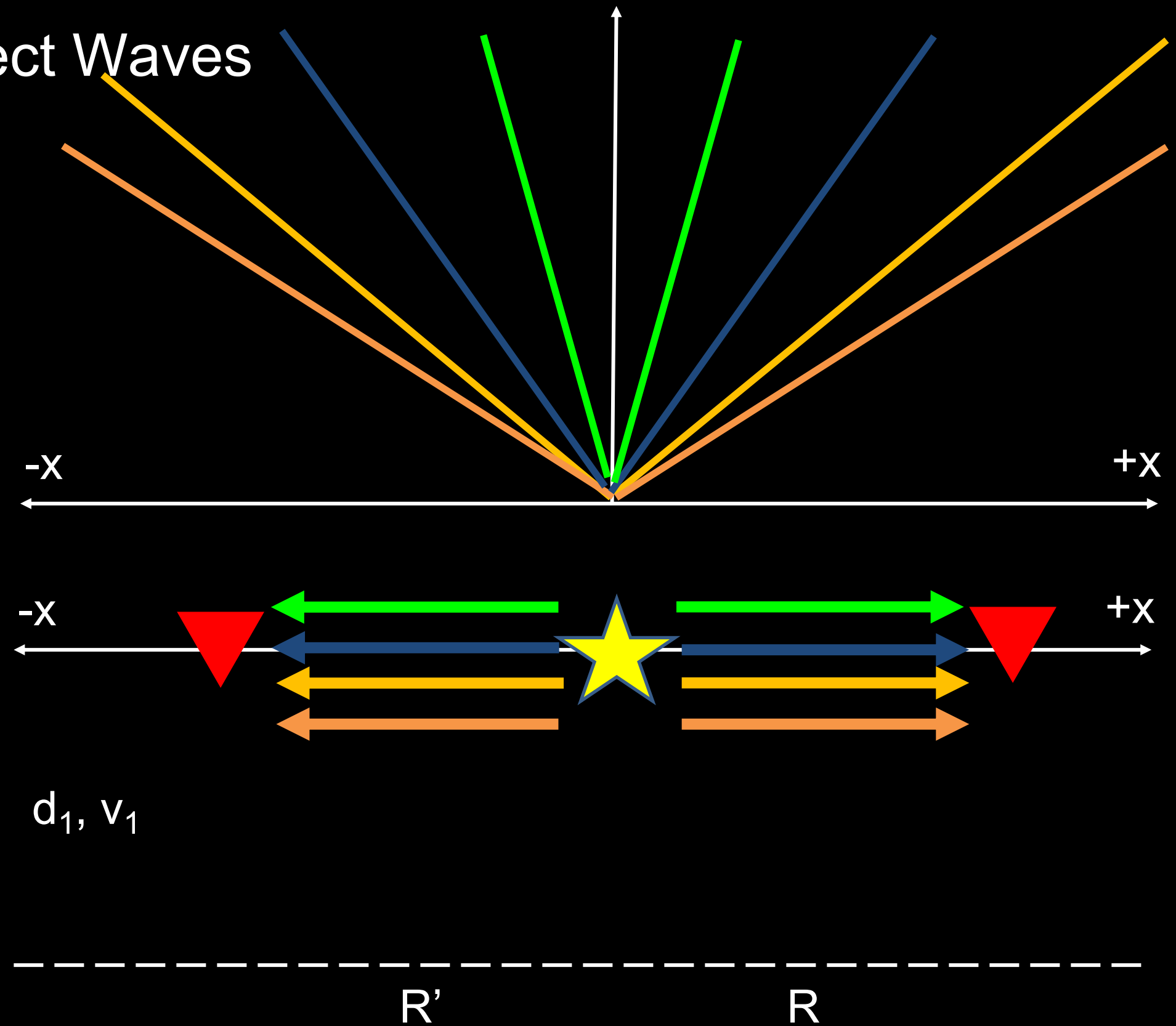


R'

R



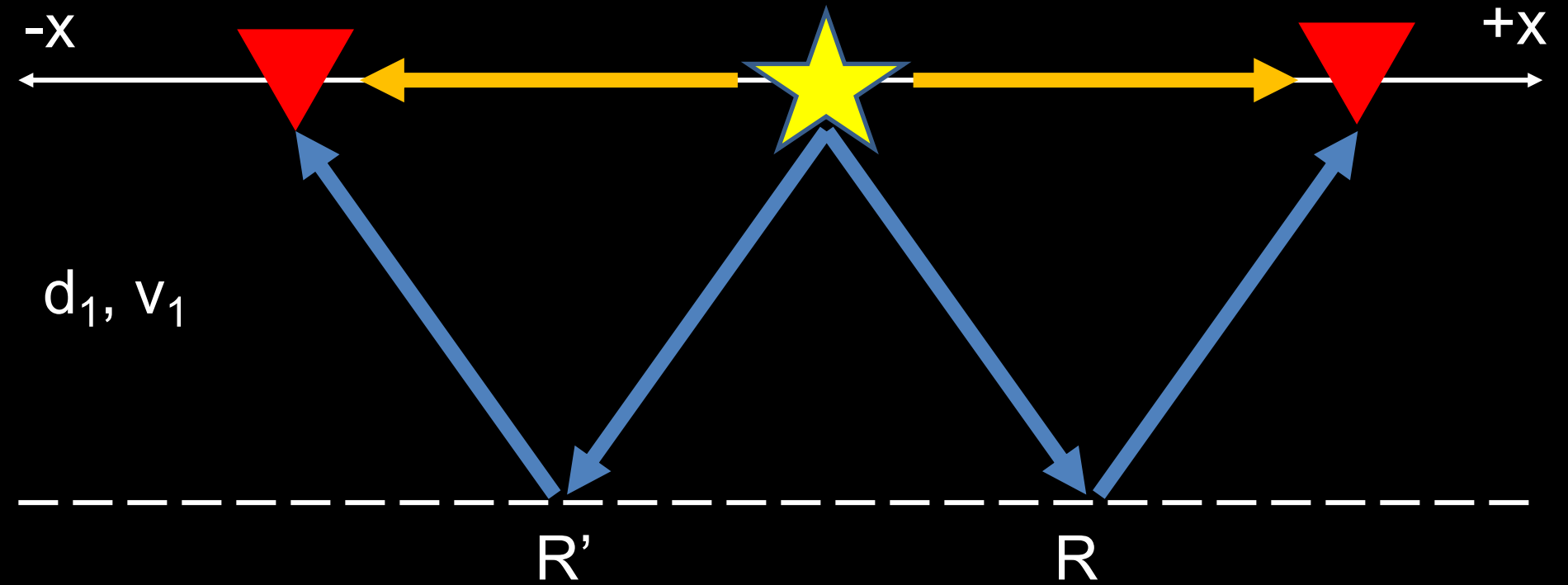
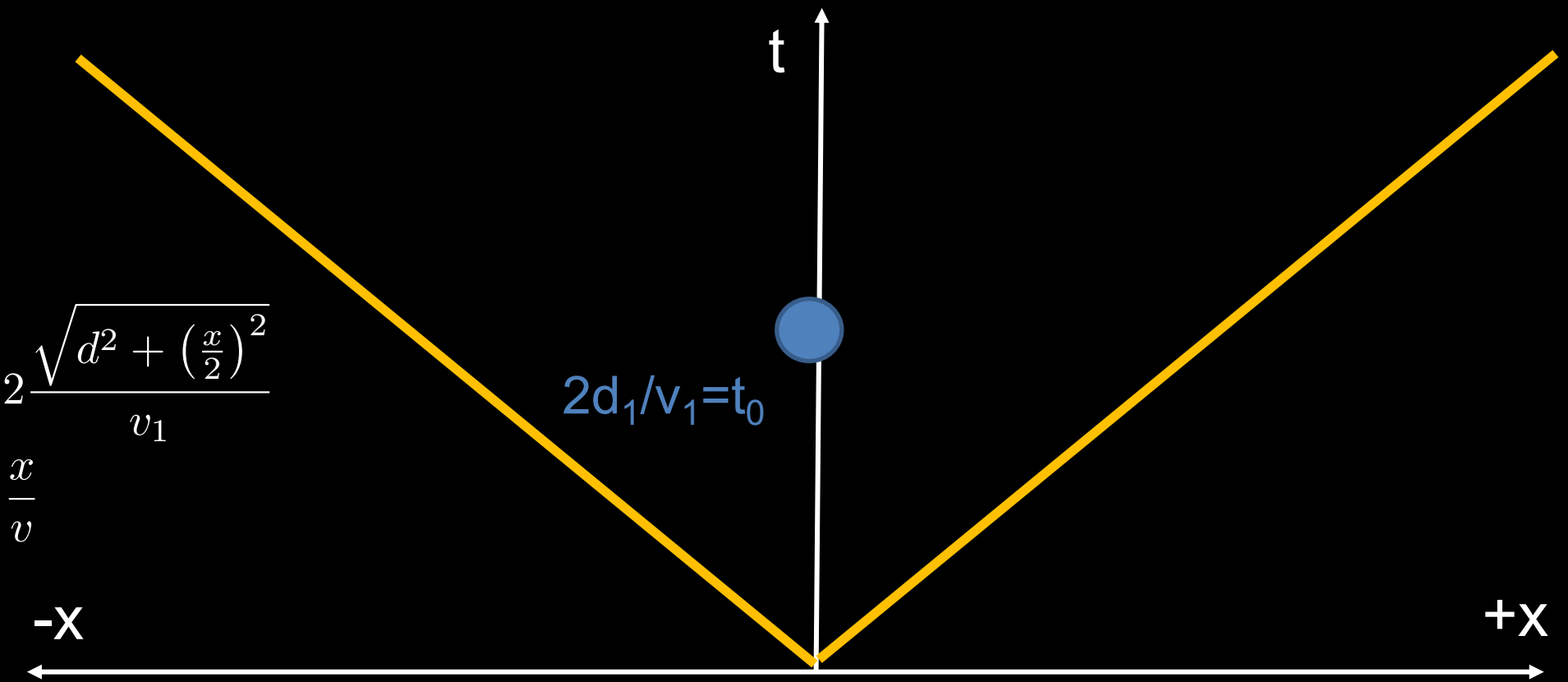
Direct Waves





$$t_r(x) = 2 \frac{\sqrt{d^2 + \left(\frac{x}{2}\right)^2}}{v_1}$$

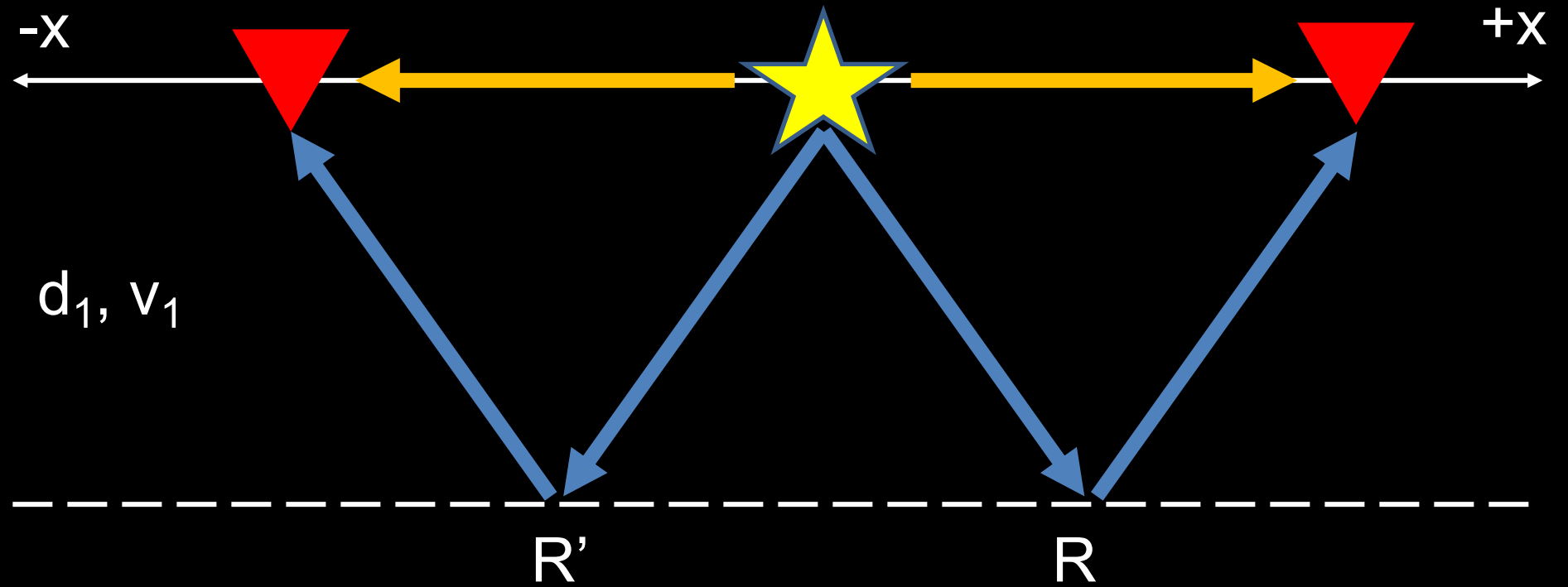
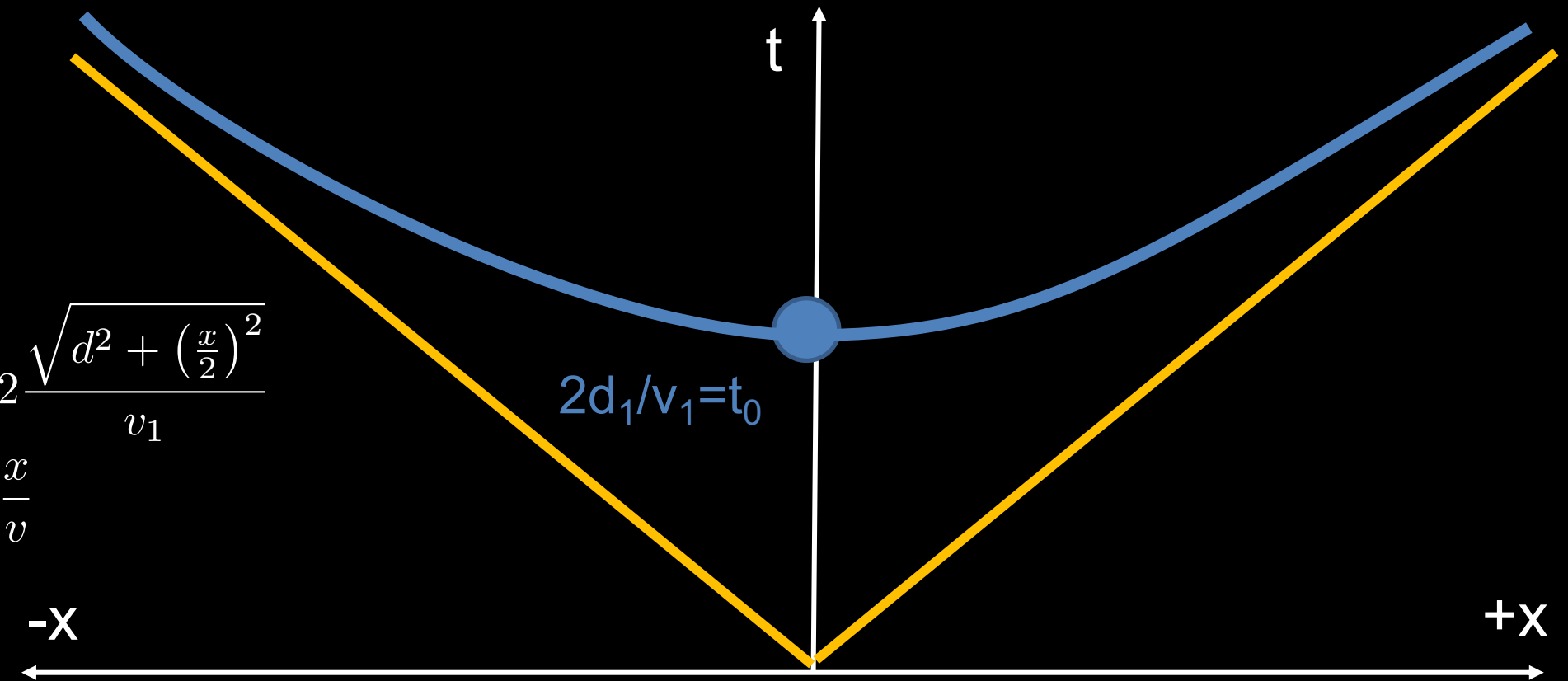
$$t_d(x) = \frac{x}{v}$$

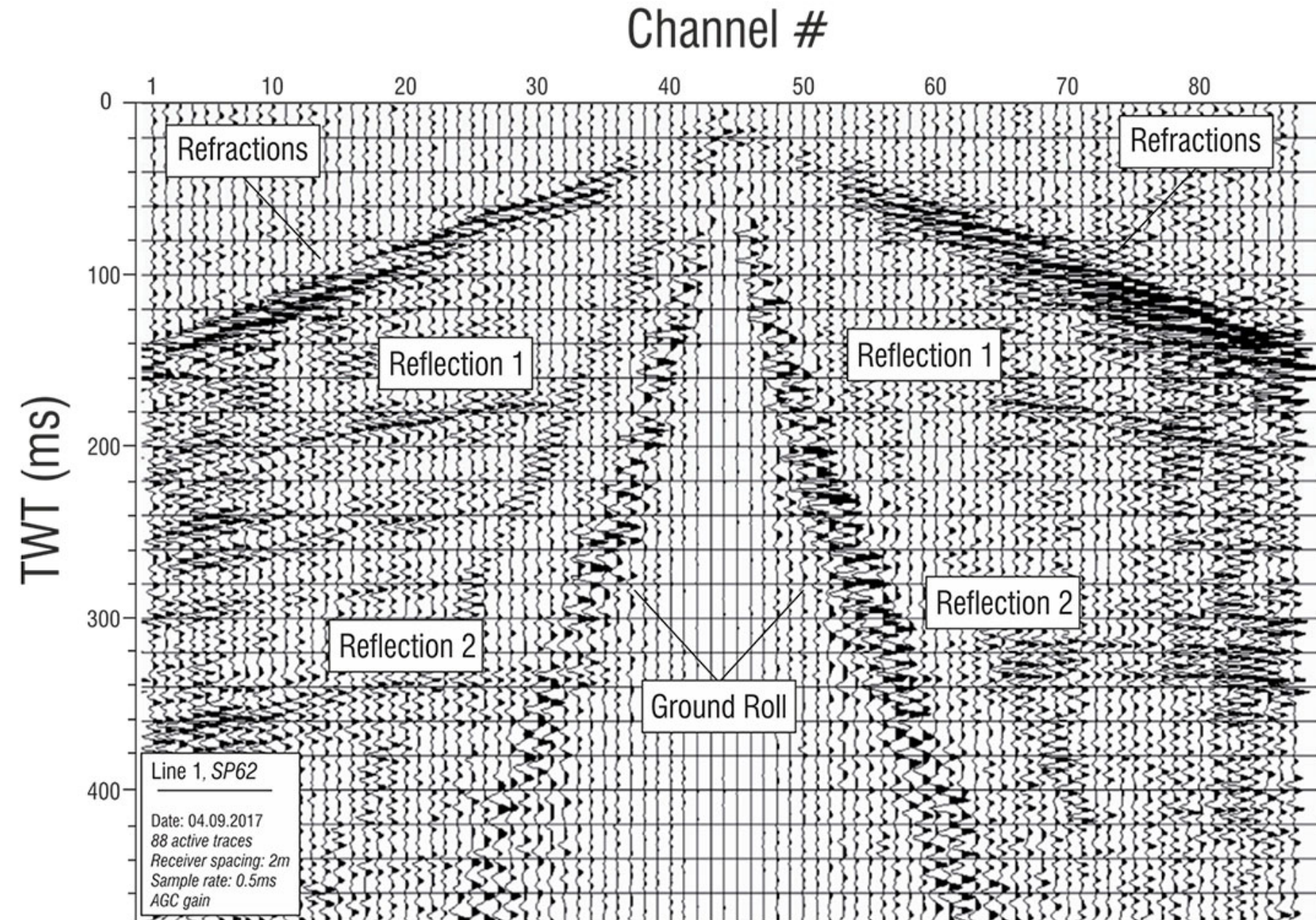


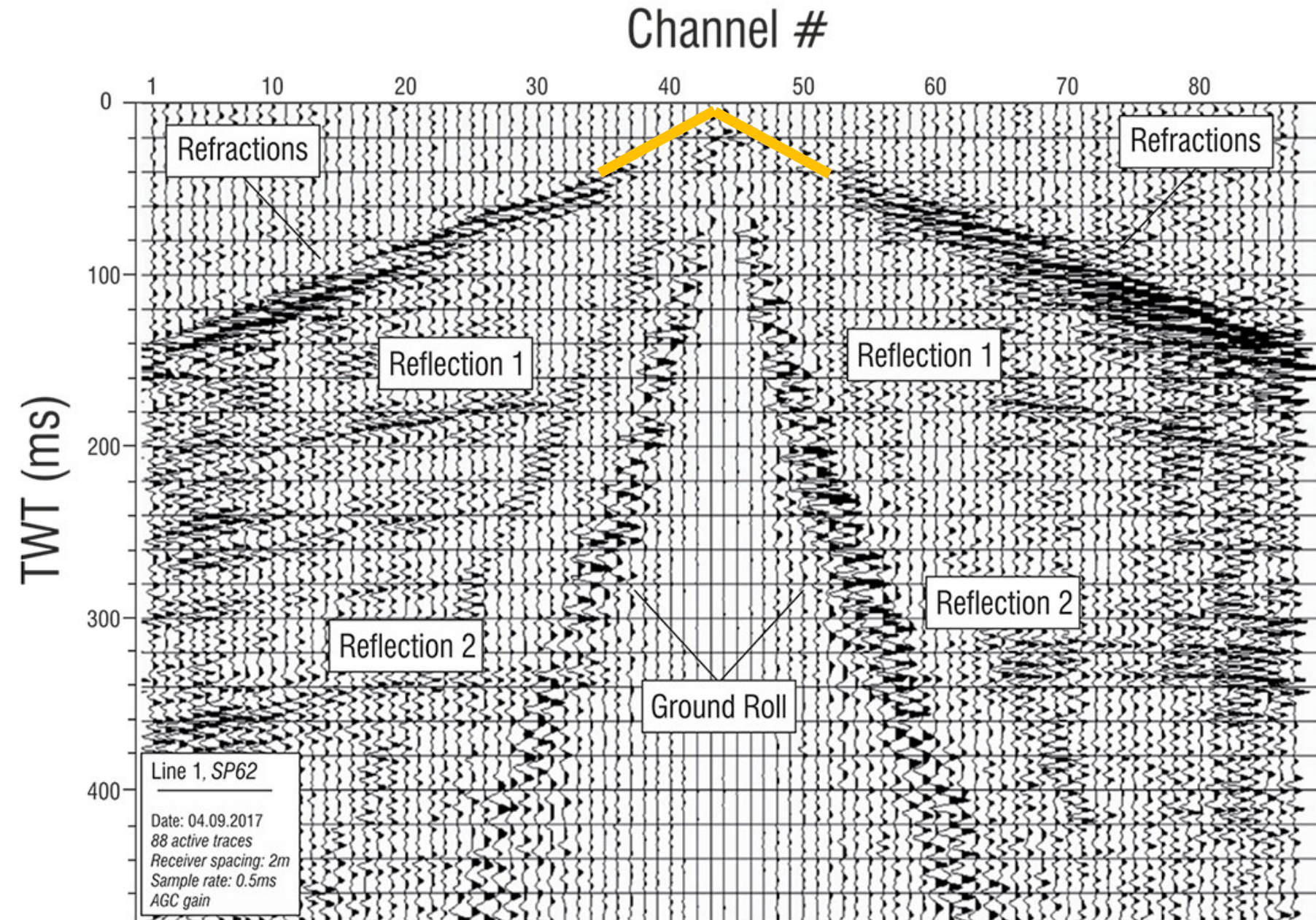
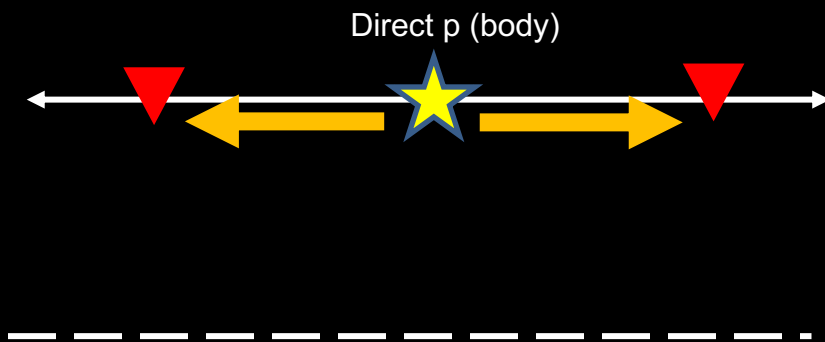


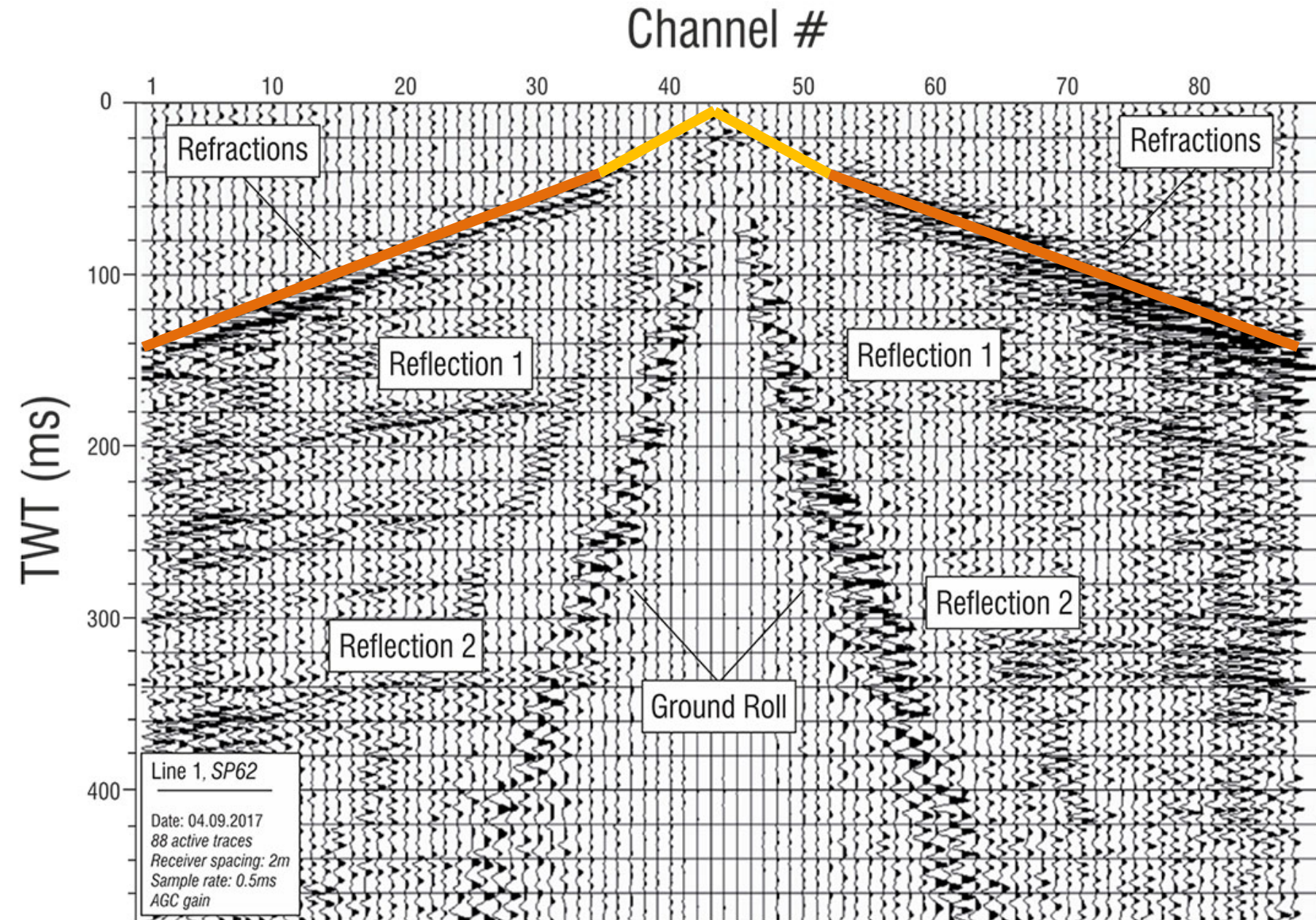
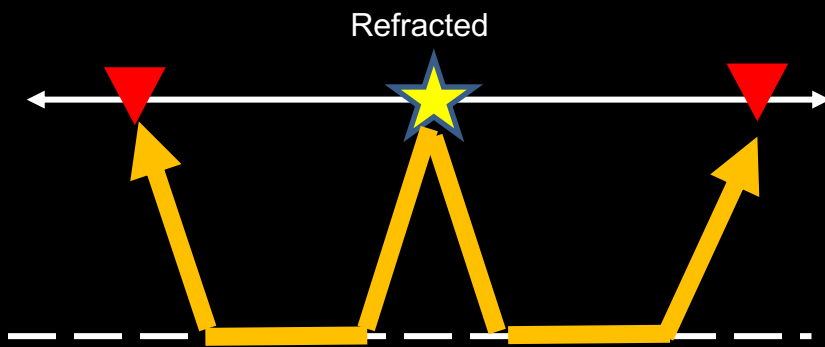
$$t_r(x) = 2 \frac{\sqrt{d^2 + \left(\frac{x}{2}\right)^2}}{v_1}$$

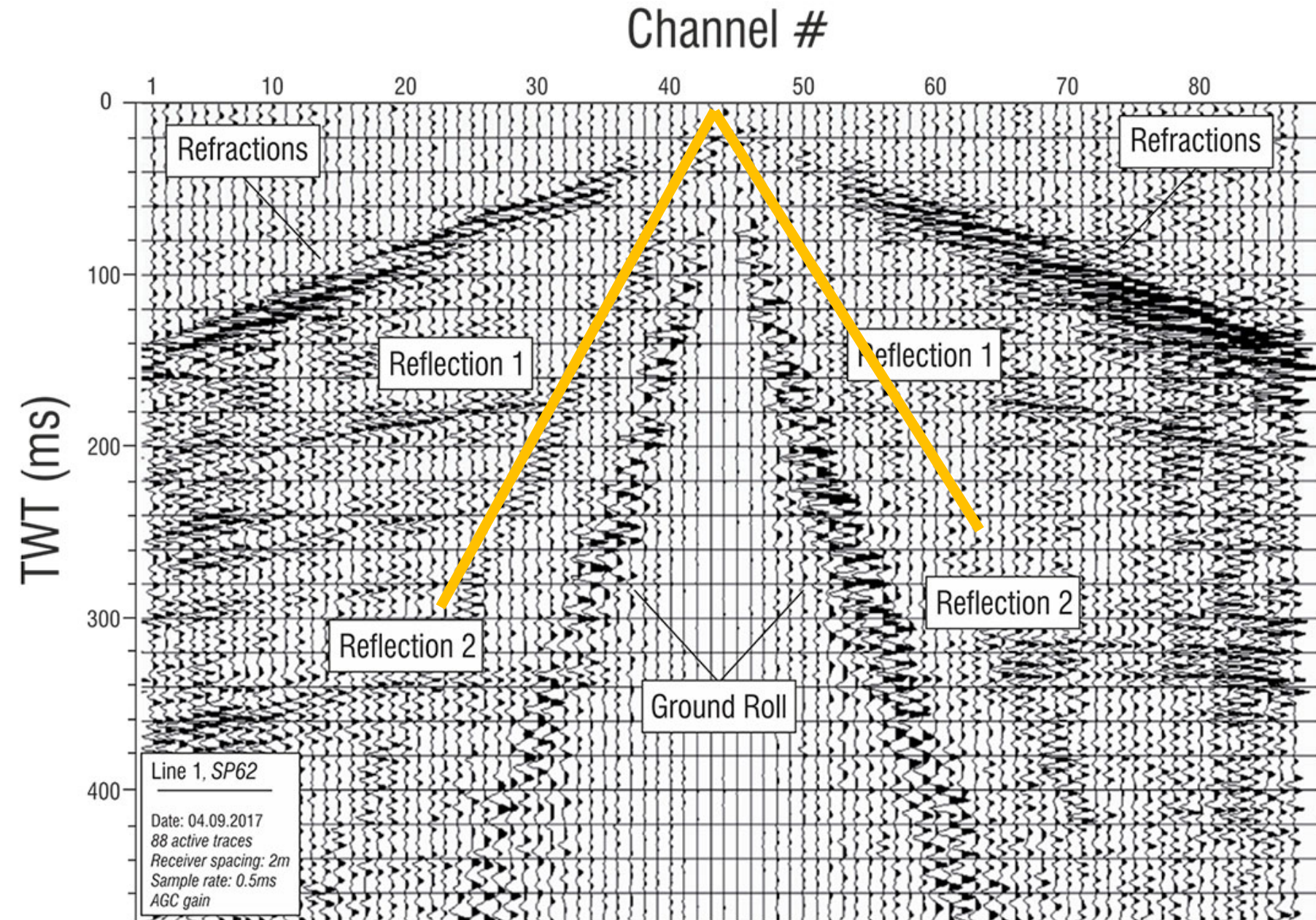
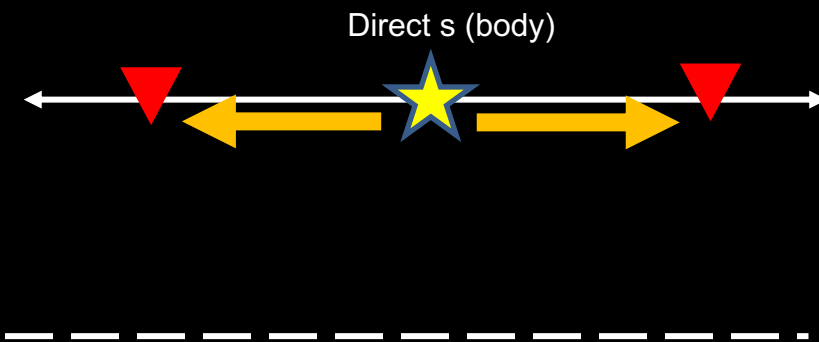
$$t_d(x) = \frac{x}{v}$$

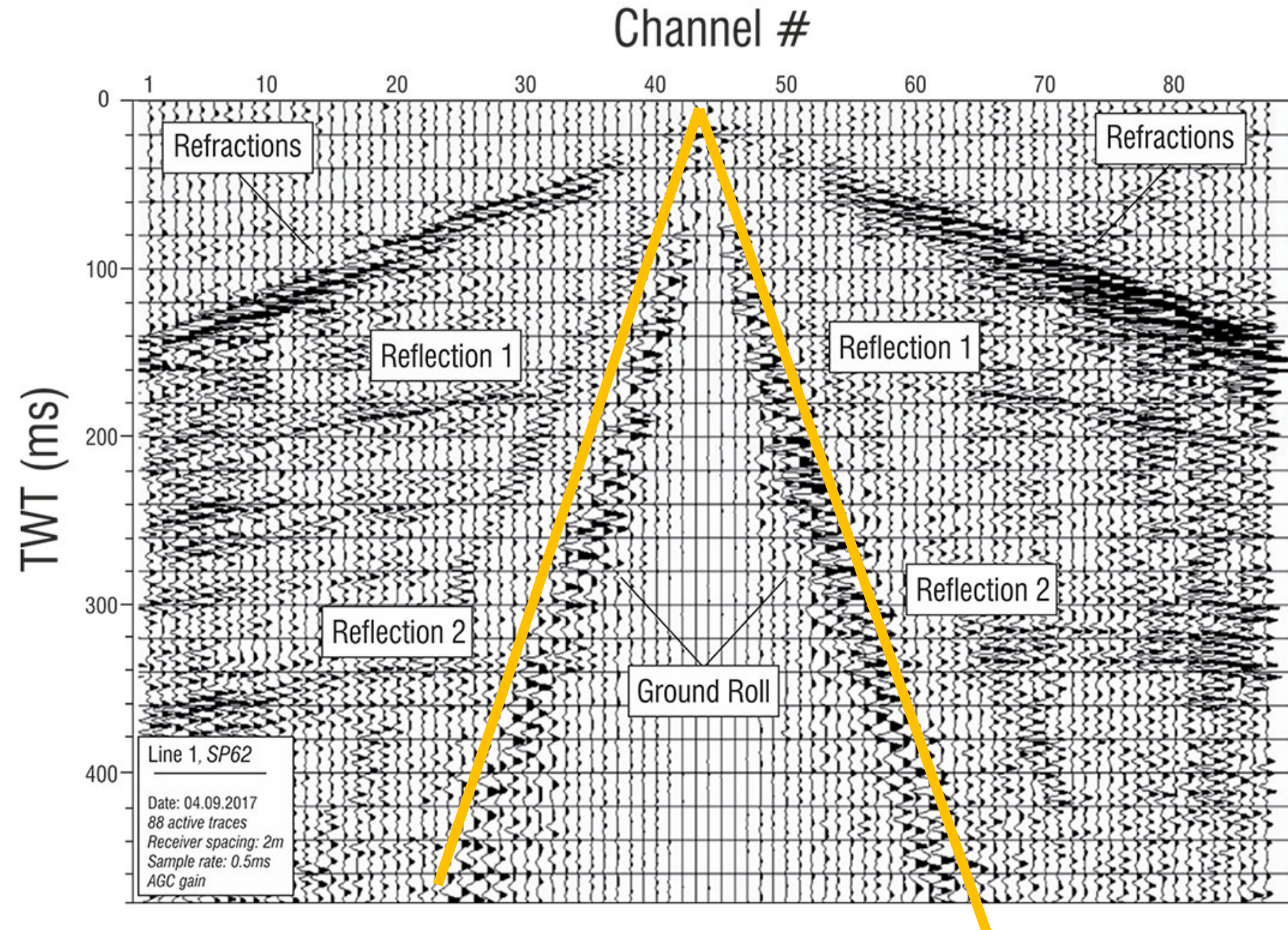
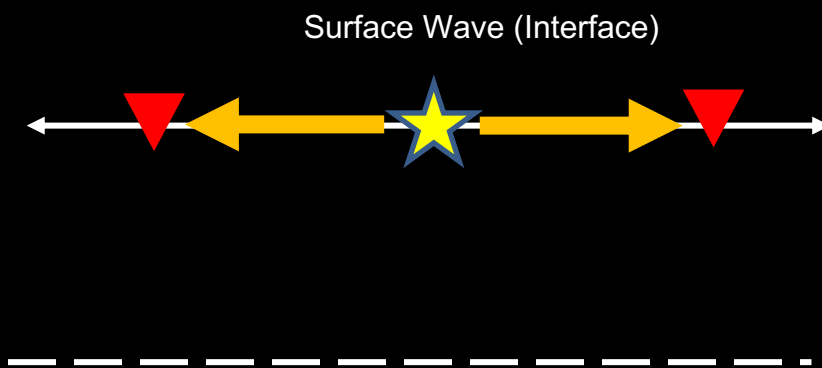


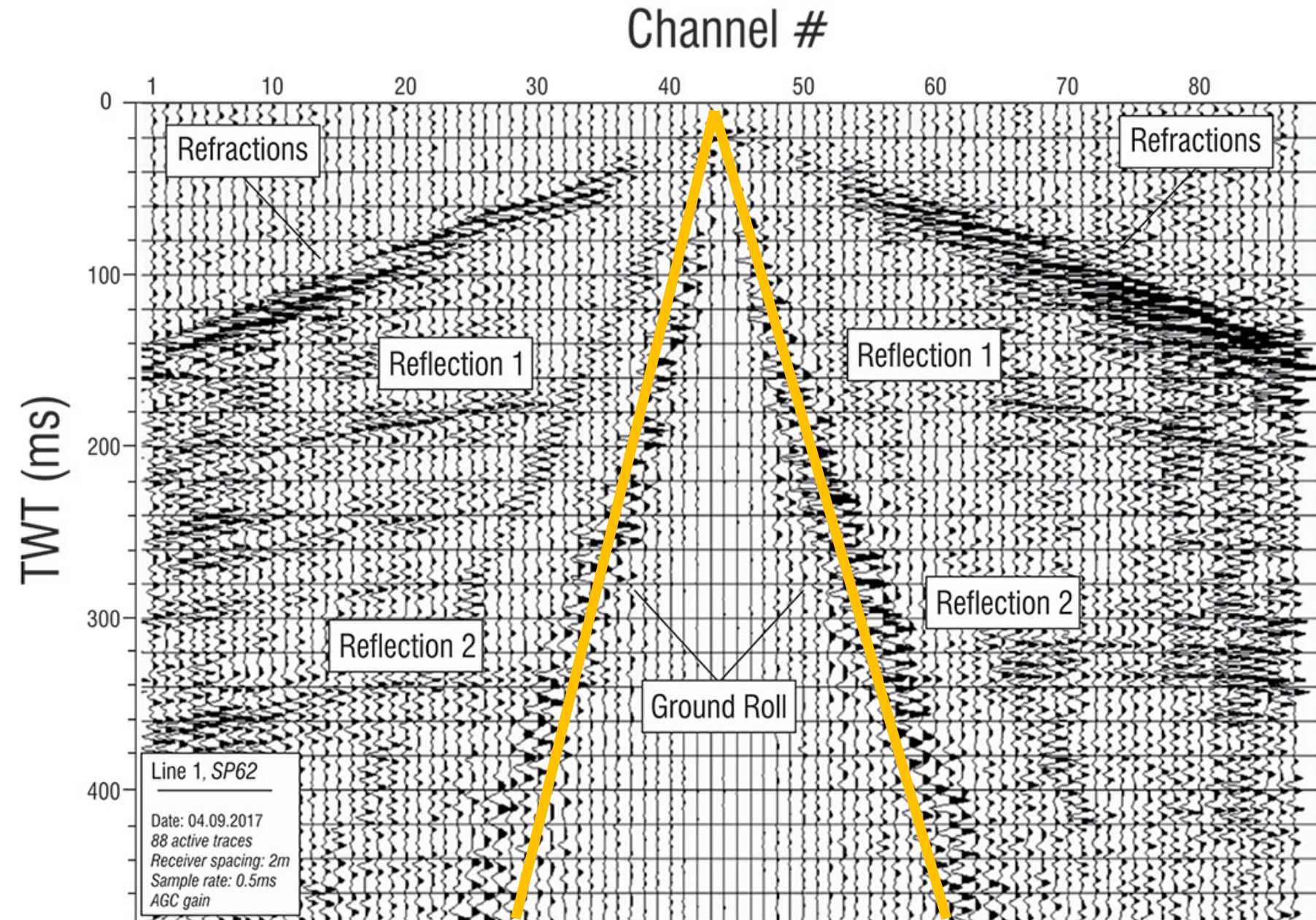
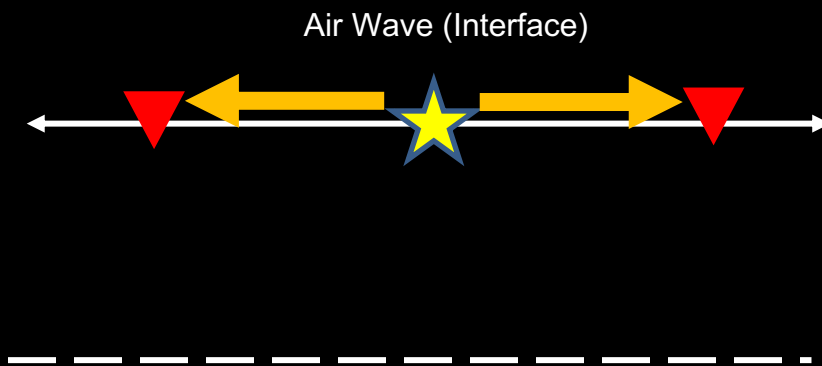


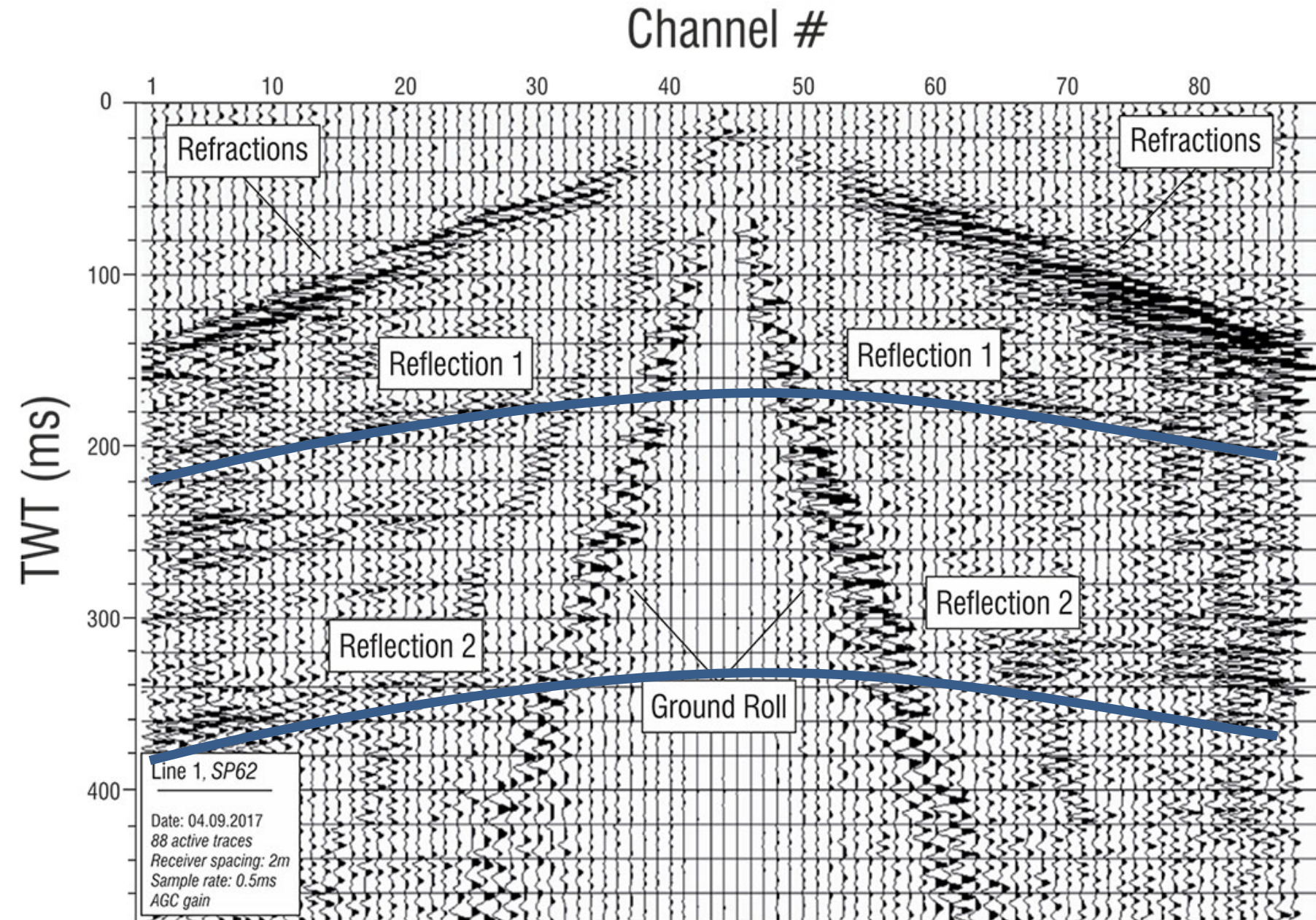
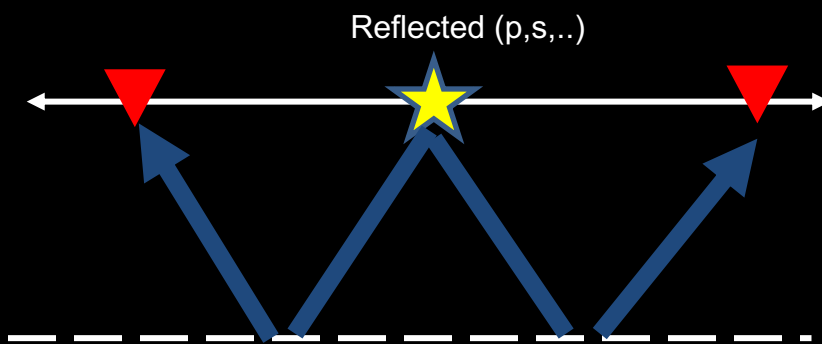














- How to determine velocities from a shot gather ?
- How to most efficiently combine different shot gathers ?



t^2-x^2 method

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$$t_r(x) = 2 \frac{\sqrt{d^2 + \left(\frac{x}{2}\right)^2}}{v_1}$$



t^2 - x^2 method

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$$t_r(x) = 2 \frac{\sqrt{d^2 + \left(\frac{x}{2}\right)^2}}{v_1}$$

$$t_r(x)^2 = \frac{4}{v_1^2} \left(d^2 + \frac{x^2}{4} \right)$$



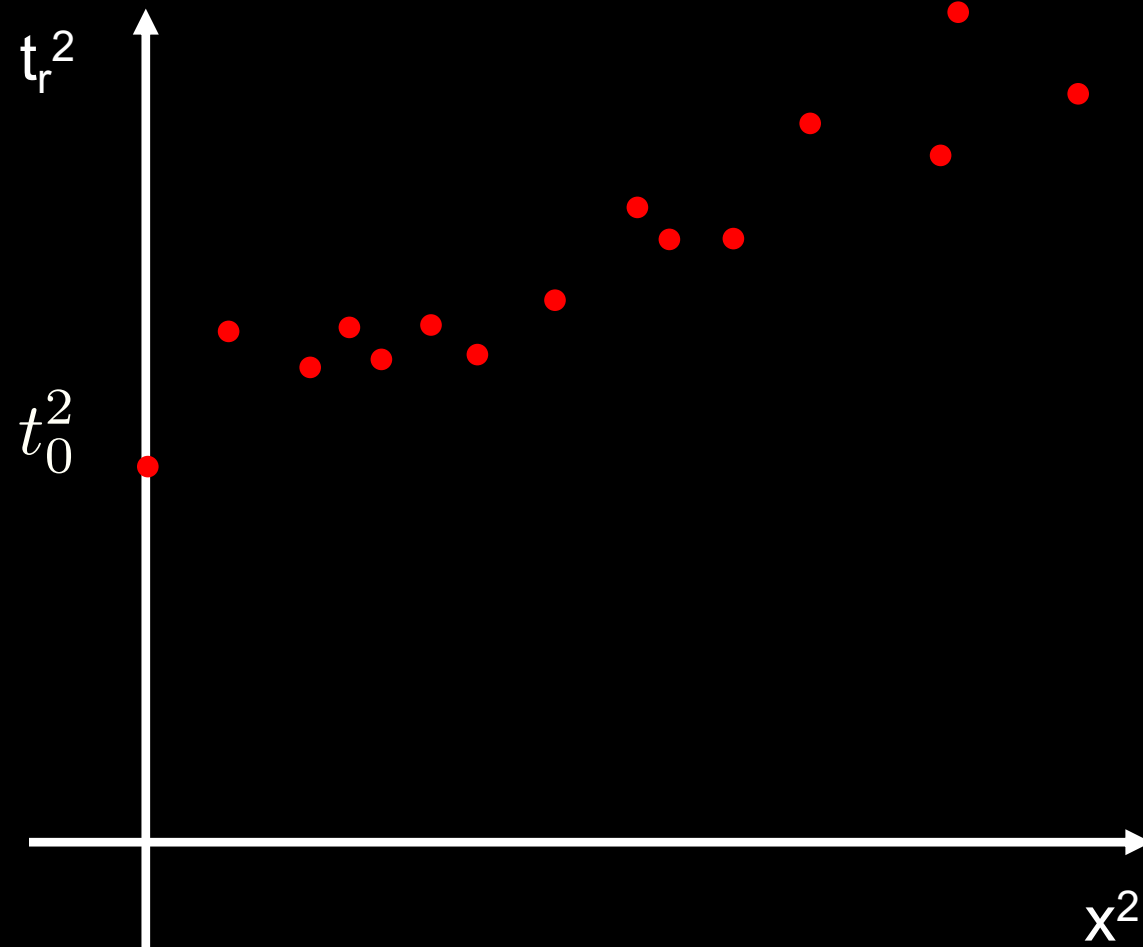
$$t_r(x) = 2 \frac{\sqrt{d^2 + \left(\frac{x}{2}\right)^2}}{v_1}$$

$$\begin{aligned} t_r(x)^2 &= \frac{4}{v_1^2} \left(d^2 + \frac{x^2}{4} \right) \\ &= \frac{1}{v_1^2} x^2 + t_0^2 \end{aligned}$$



t^2 - x^2 method

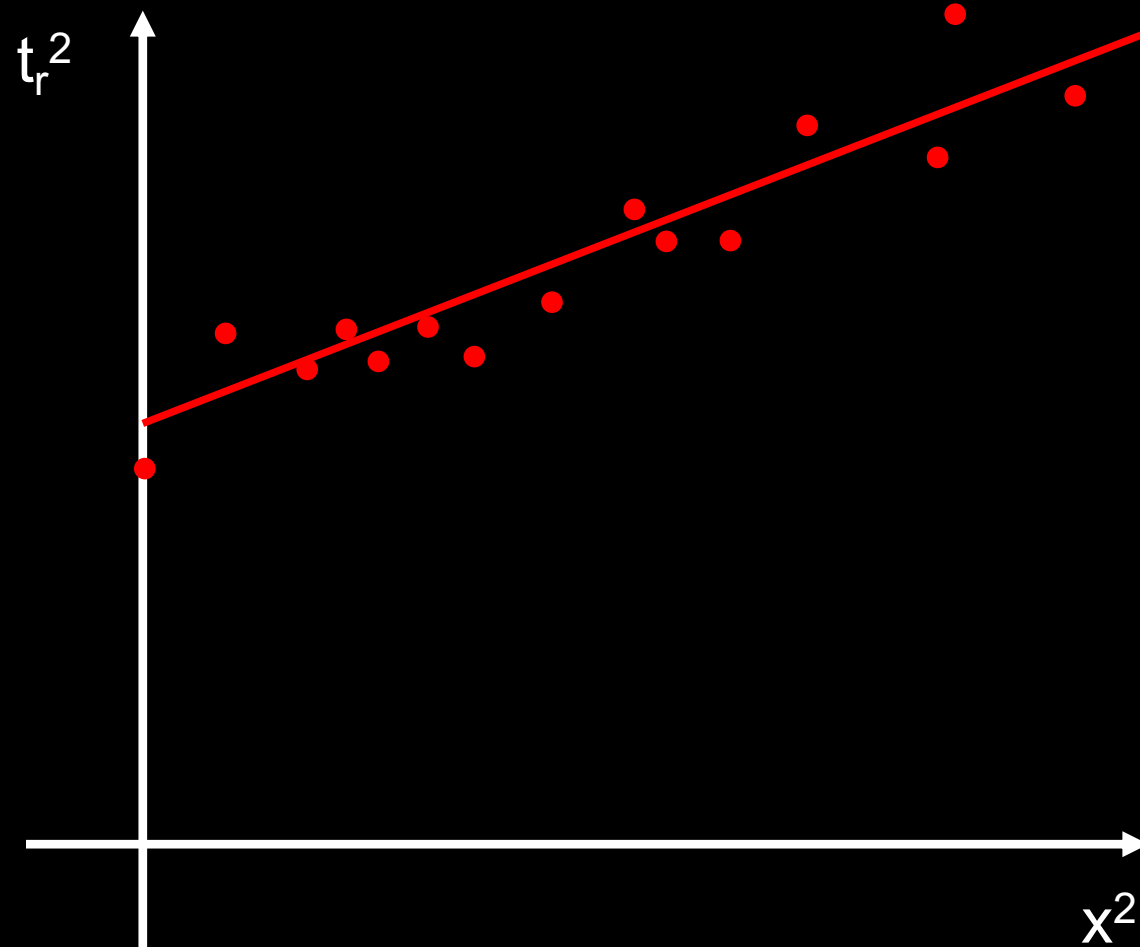
$$t_r(x)^2 = \frac{1}{v_1^2} x^2 + t_0^2$$





t^2 - x^2 method

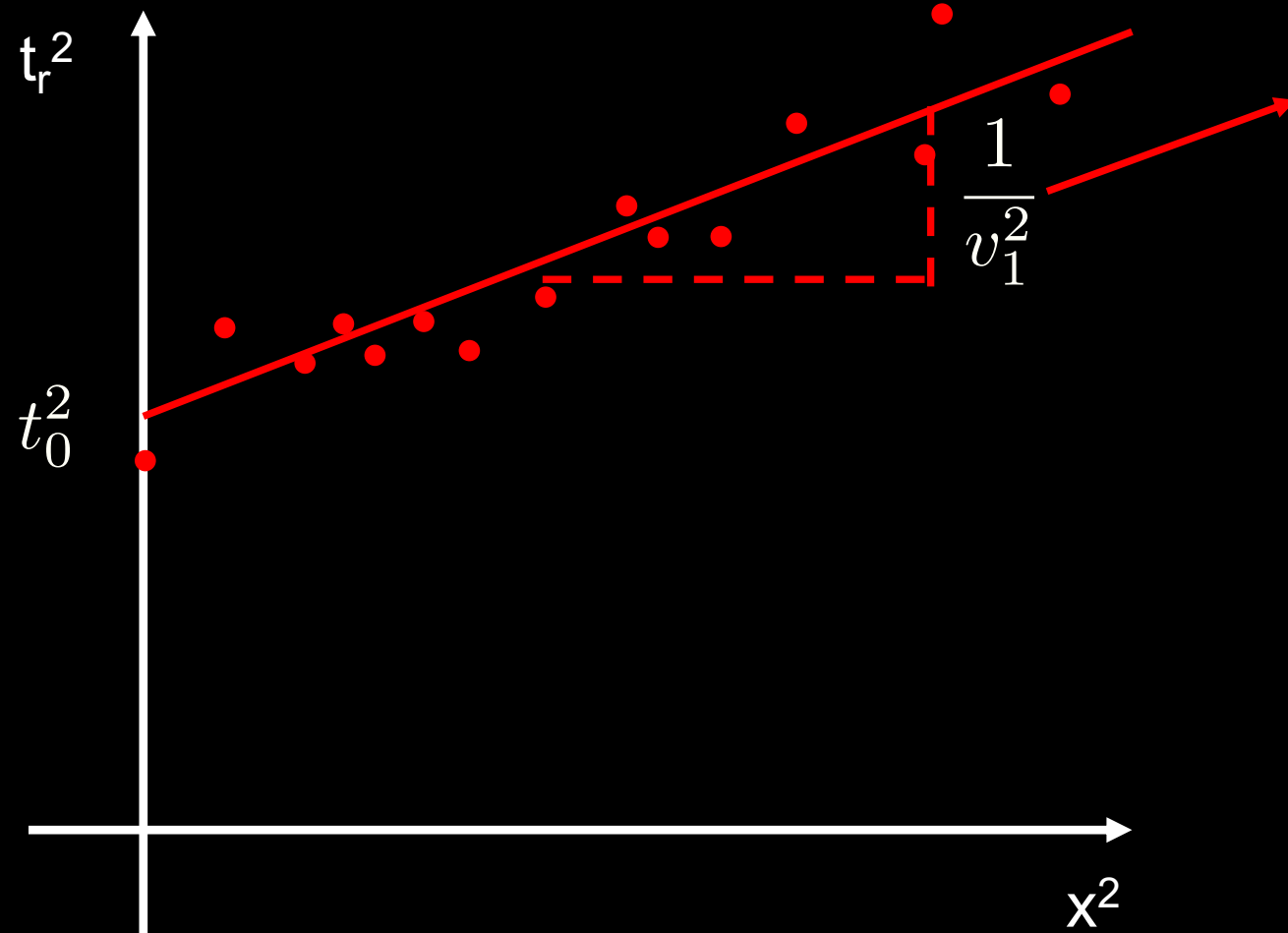
$$t_r(x)^2 = \frac{1}{v_1^2} x^2 + t_0^2$$





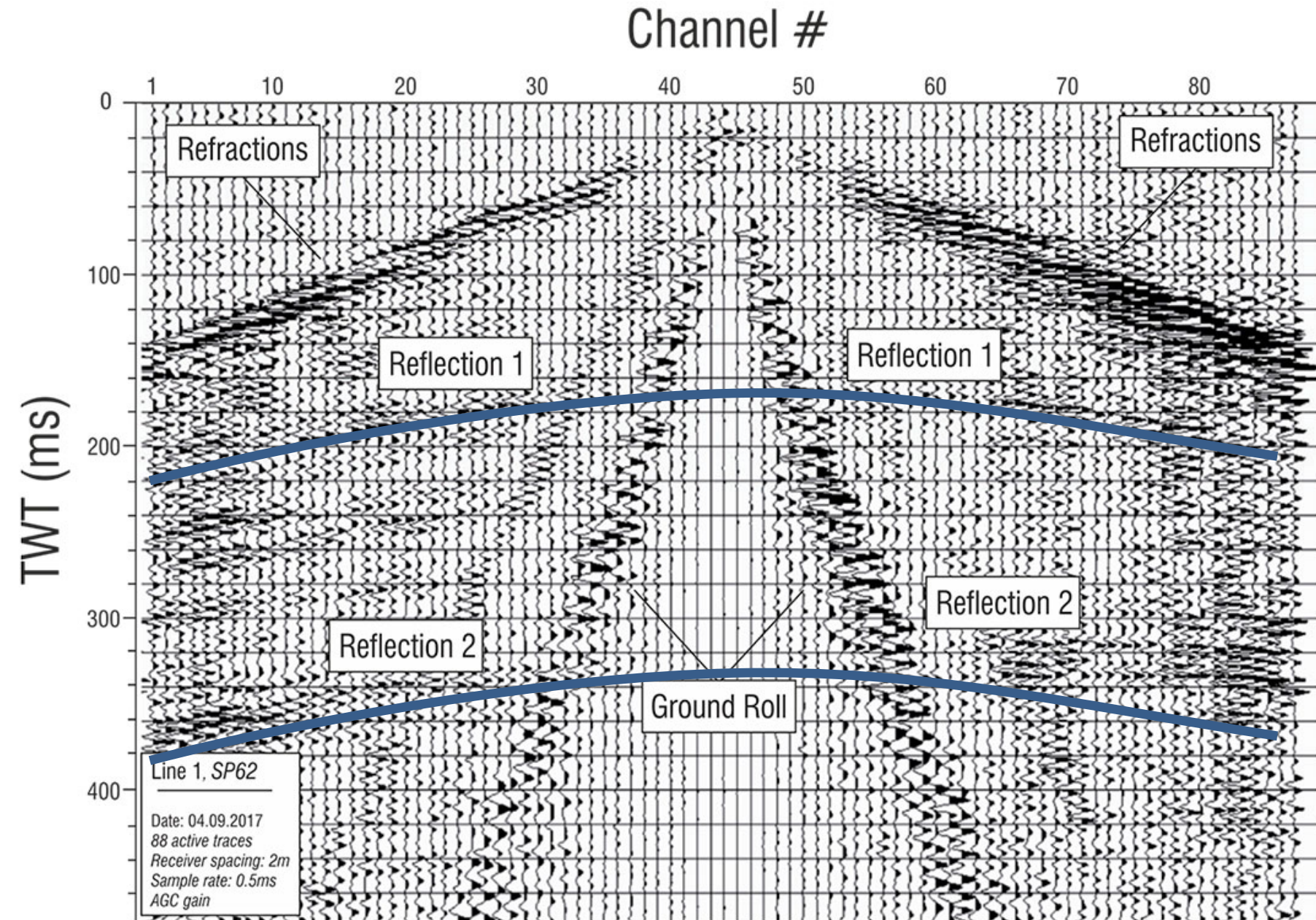
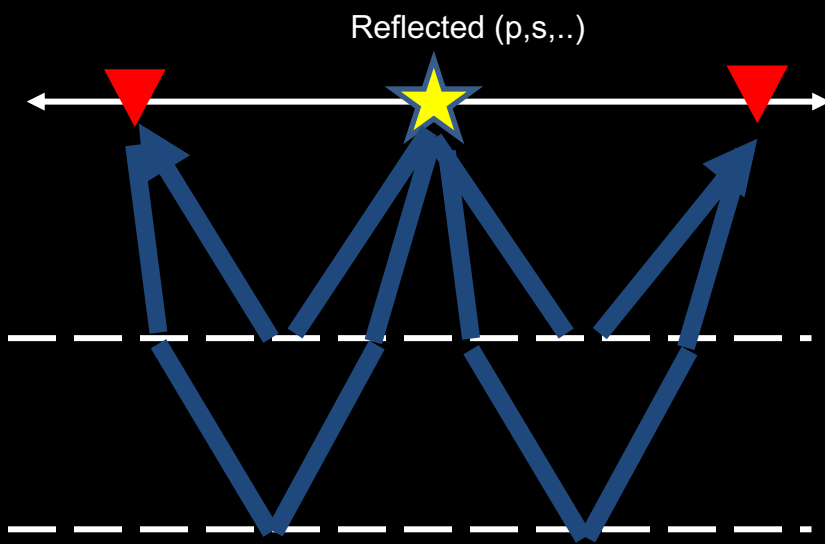
t^2 - x^2 method

$$t_r(x)^2 = \frac{1}{v_1^2} x^2 + t_0^2$$





- What about multiple layers ?

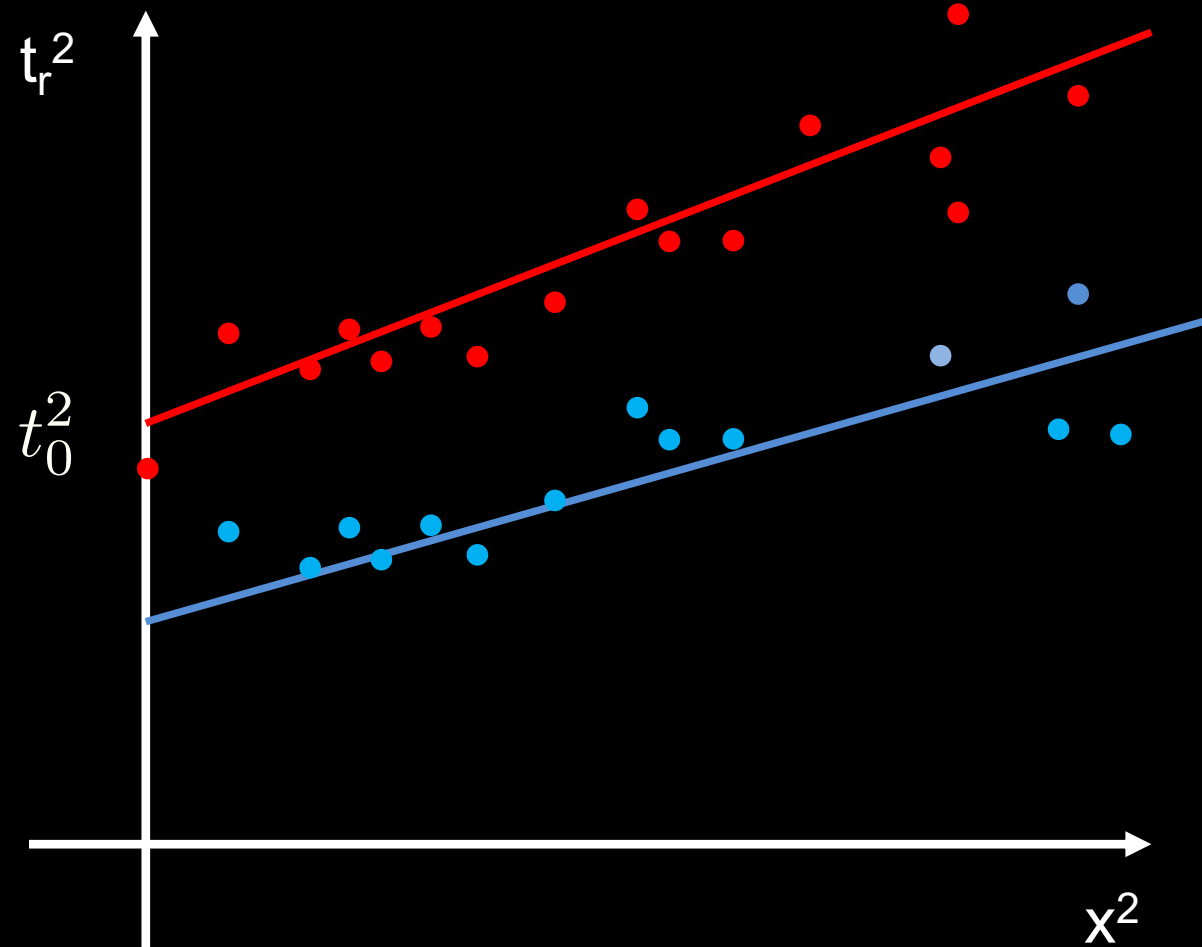




t^2 - x^2 method

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$$t_r(x)^2 = \frac{1}{v_1^2} x^2 + t_0^2$$

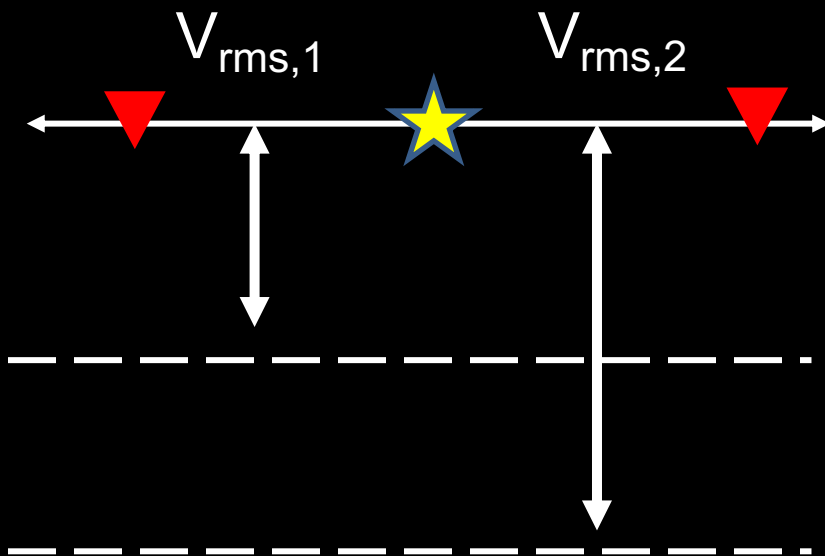


- This works for near-nadir incidence



RMS Velocities

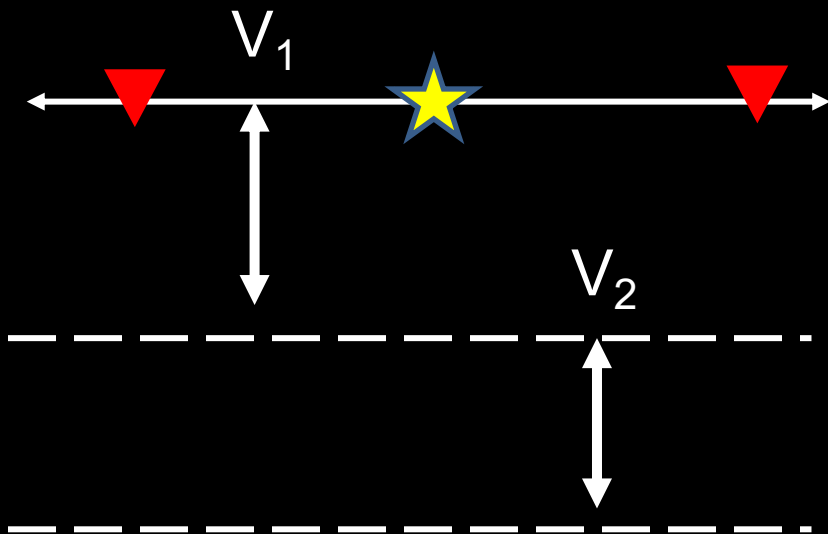
- Methods mentioned above deliver average velocities from the surface to the reflector.





Interval Velocities

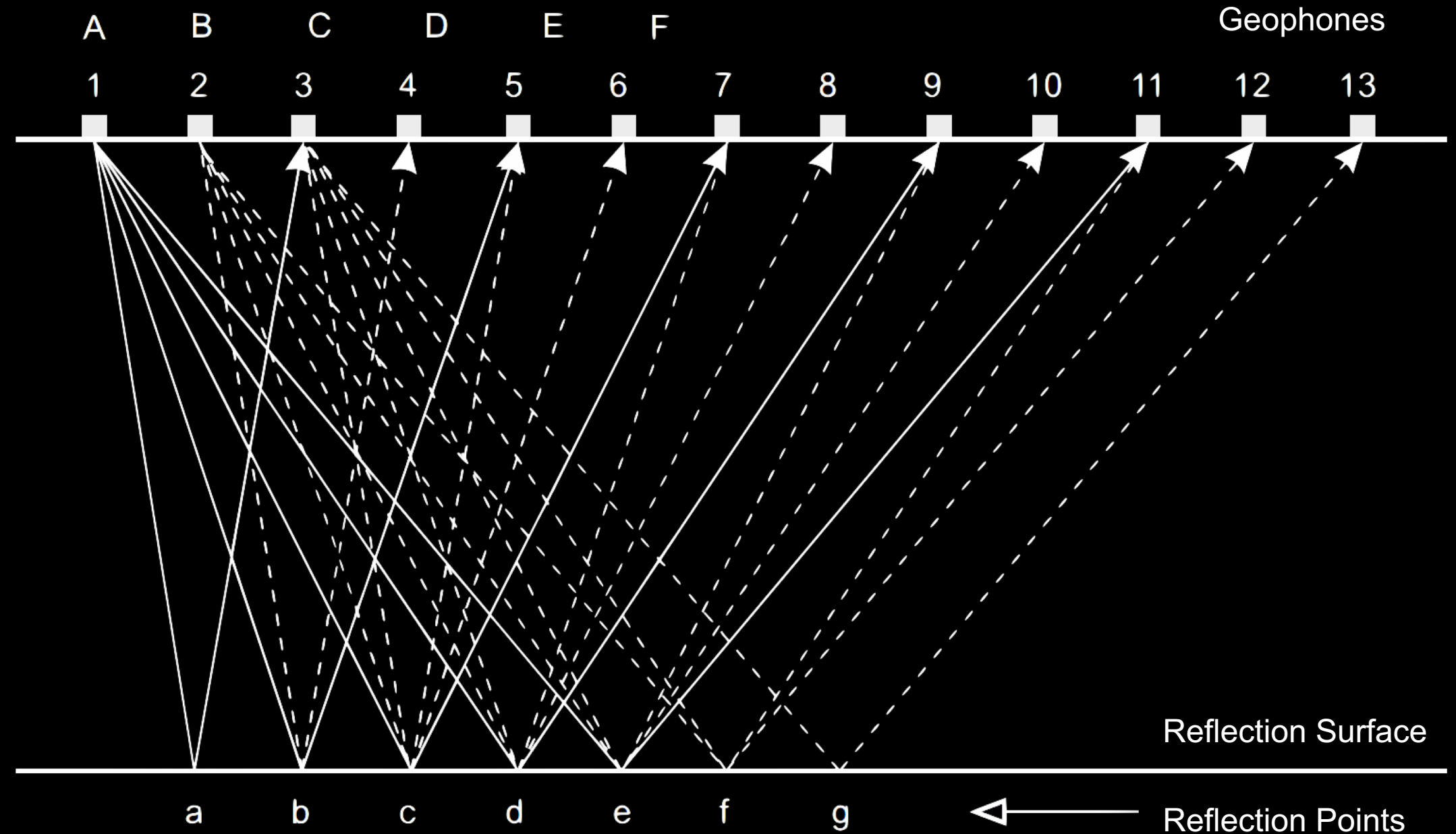
- Interval velocities can be obtained in a top-to-bottom approach (Dix Inversion)



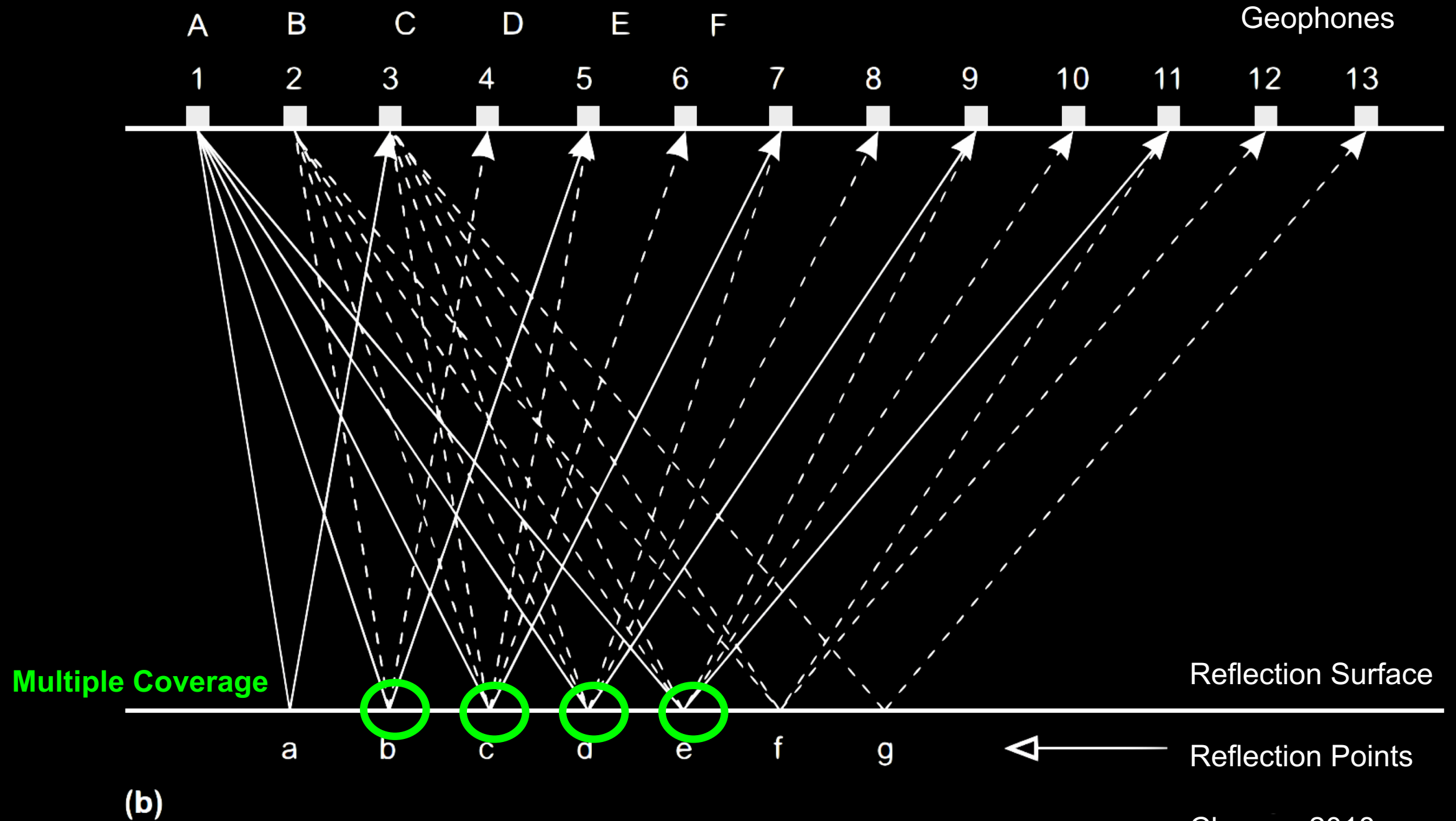


Shot Gather and CMP Gather

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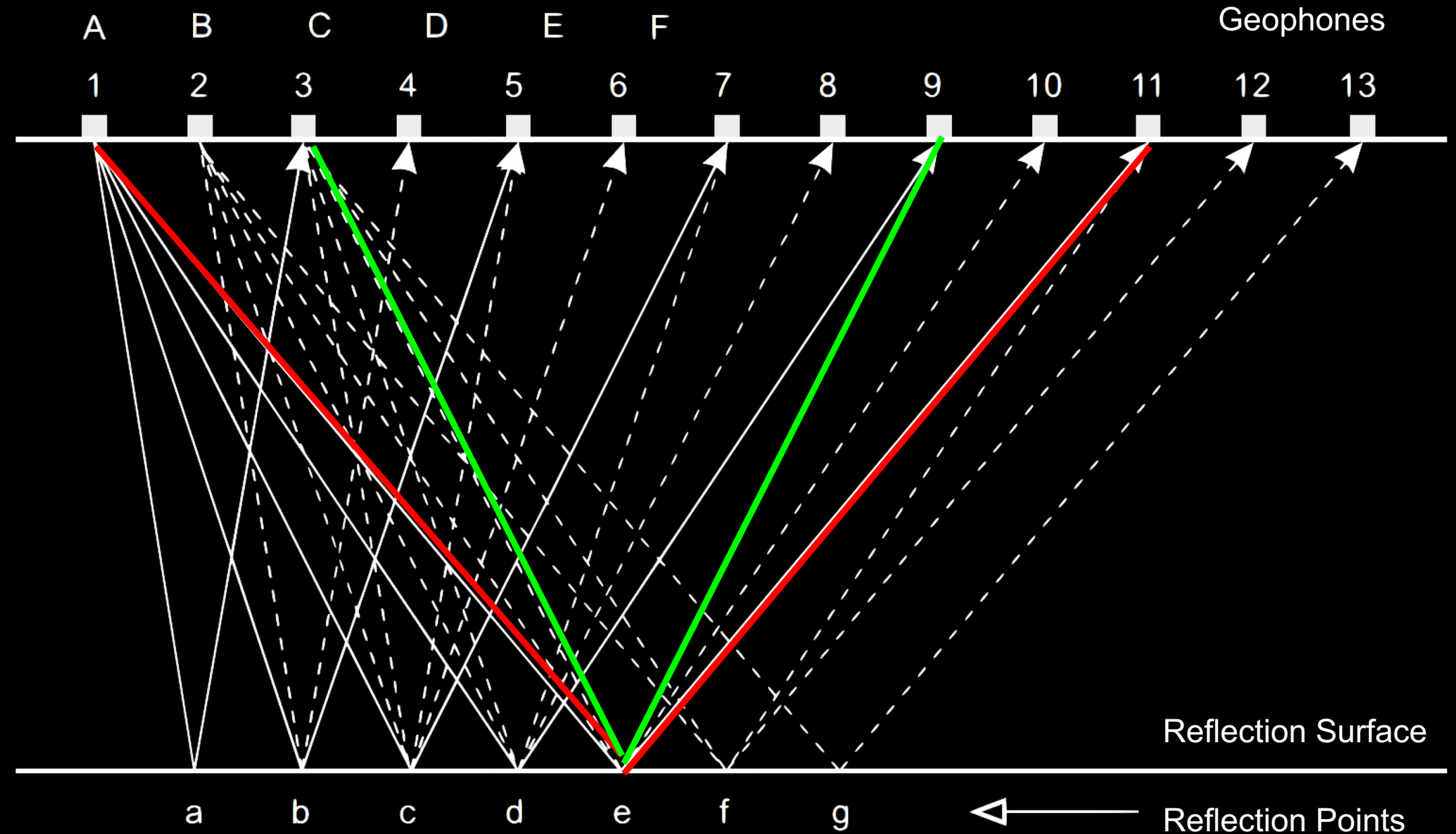
(b)





Shot Gather to CMP Gather

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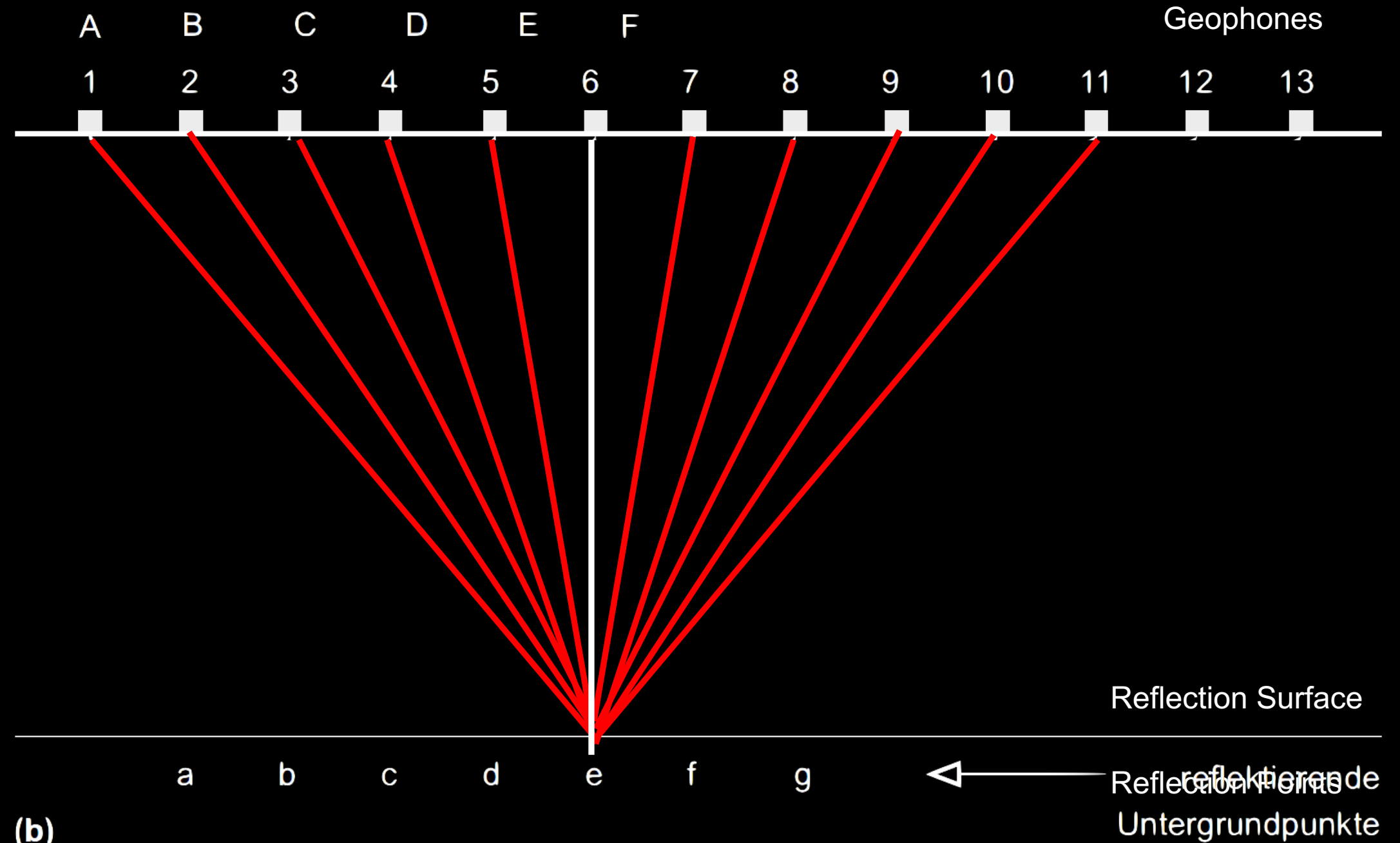


(b)



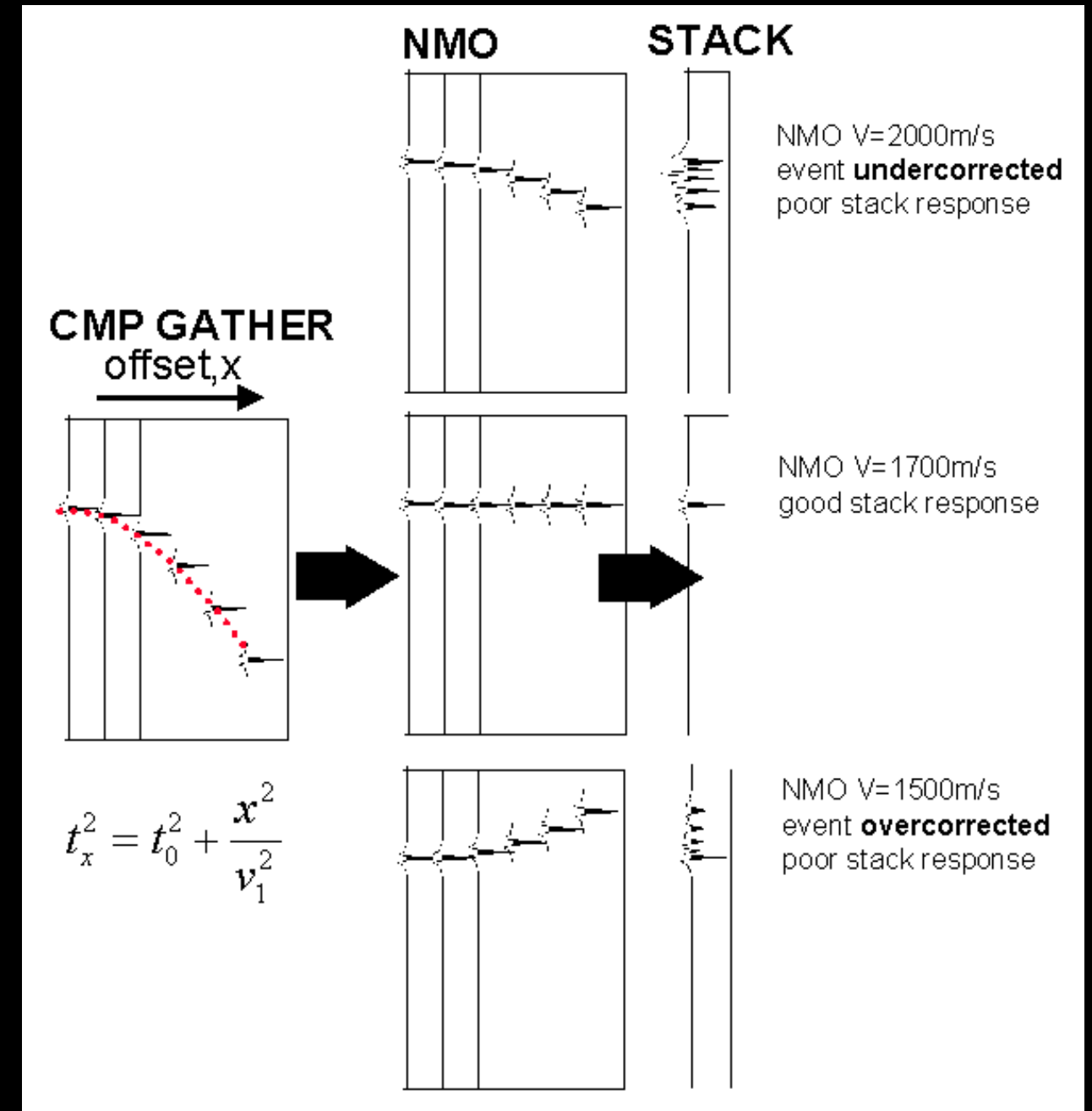
Shot Gather to CMP Gather

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- NMO stacking in CMP gather

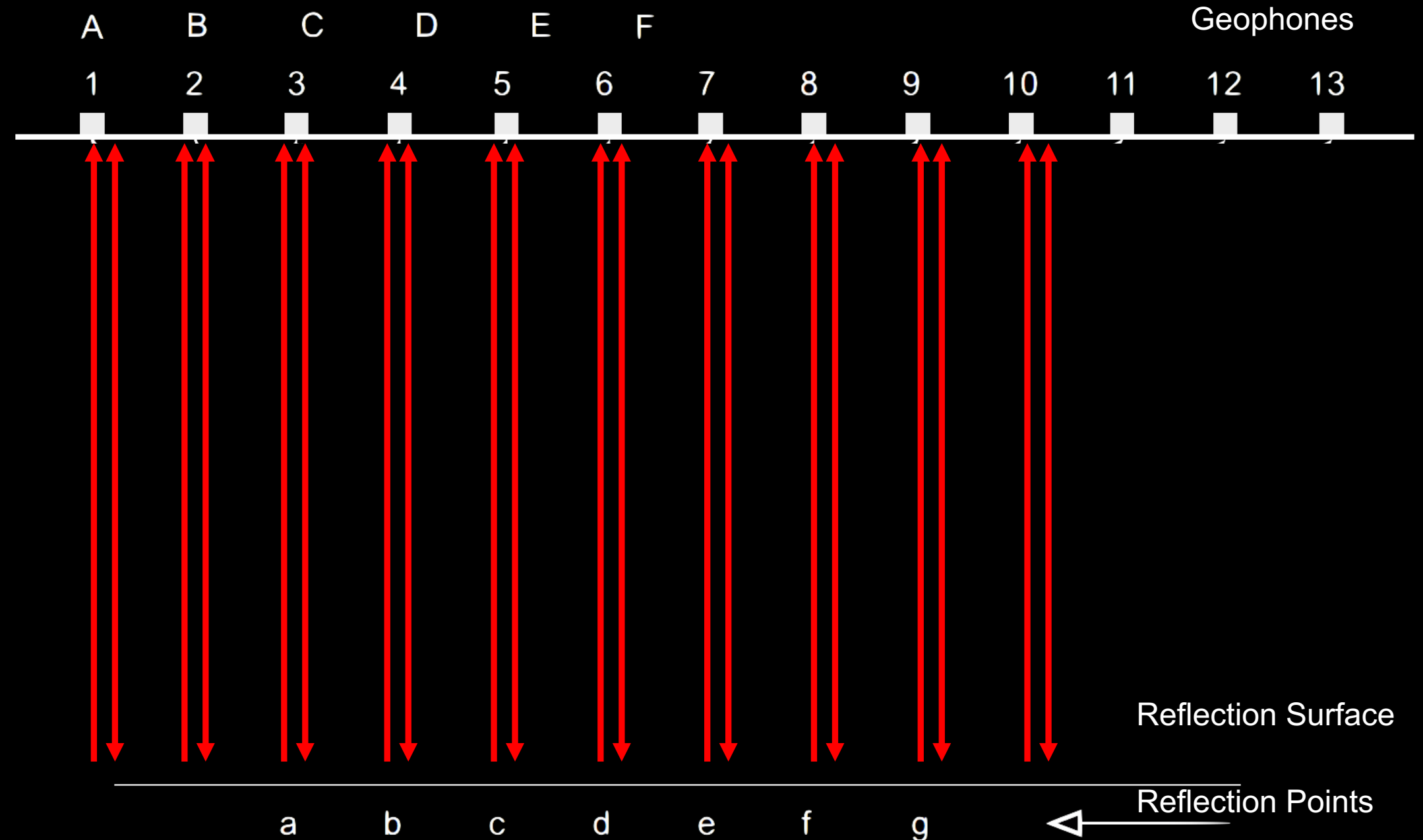




Zero Offset Gather

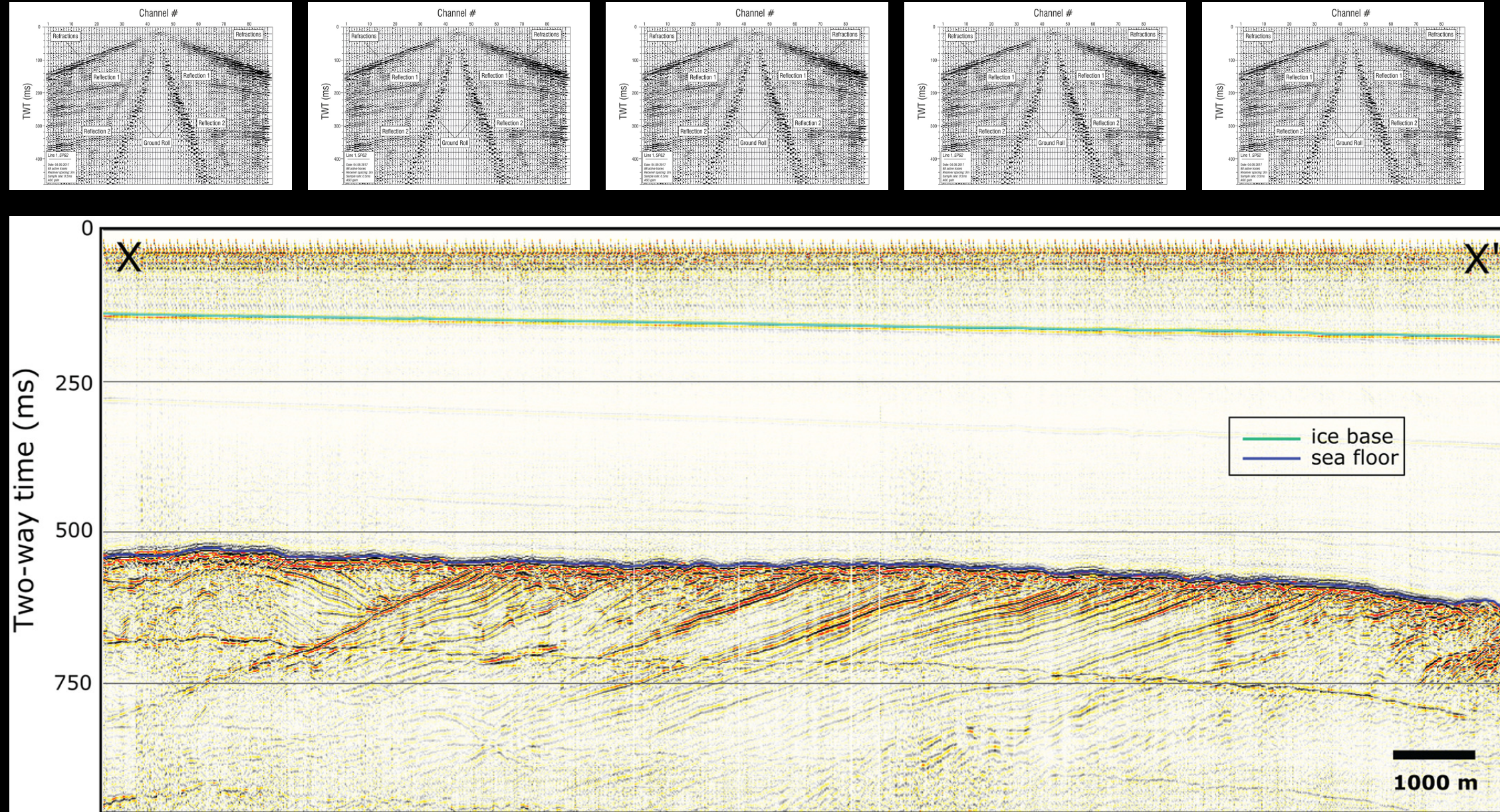
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The stacking
enables to
combine
everything into
quasi vertical
raypaths



(b)

Shot gather, CMP sorting, NMO
correction, stacking



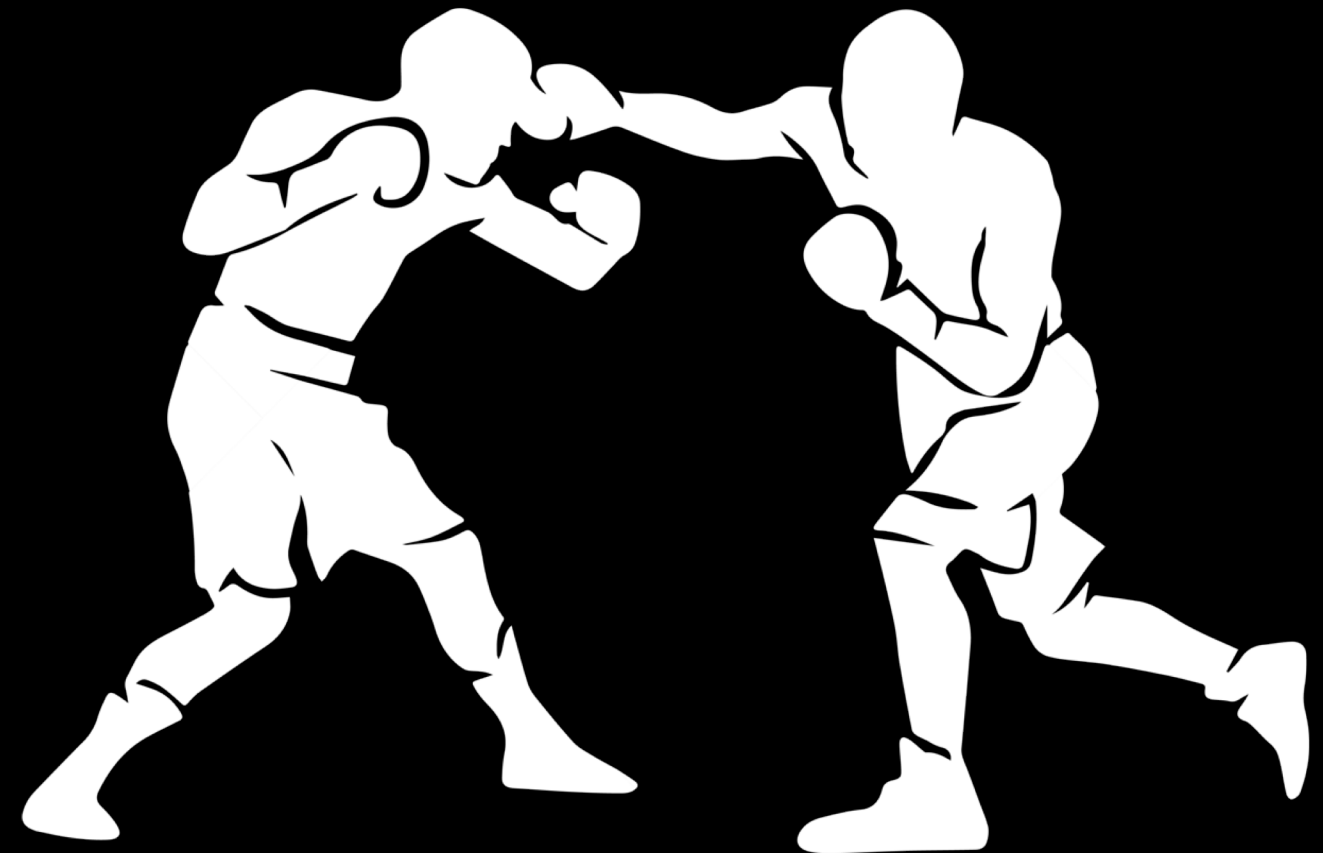


- The zero offset gather provides an “image” of the subsurface. But it is important to realize that this “image” does in general not correspond to the true geometry of the sub-surface. Why?
- Because so far we have done everything with horizontal reflectors. Off-angle point targets or dipping interfaces will have a different appearance in (unmigrated) reflection datasets.



Refraction

Reflection



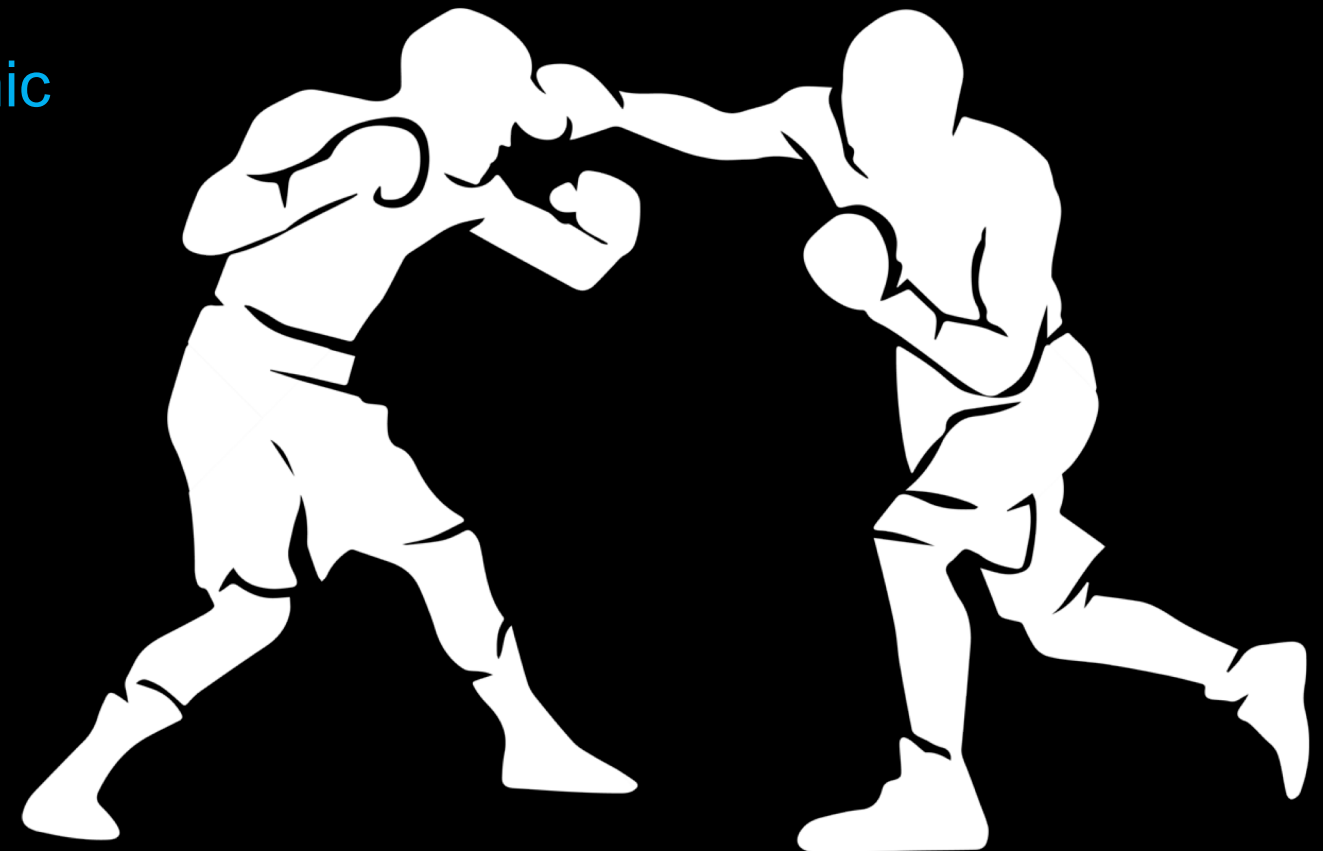
[academic use]



- + No need for critical angles
- + No need for increasing seismic velocities with depth
- + Better suited for deep exploration
- + Less restrictions on sub-surface geometry

Refraction

Reflection



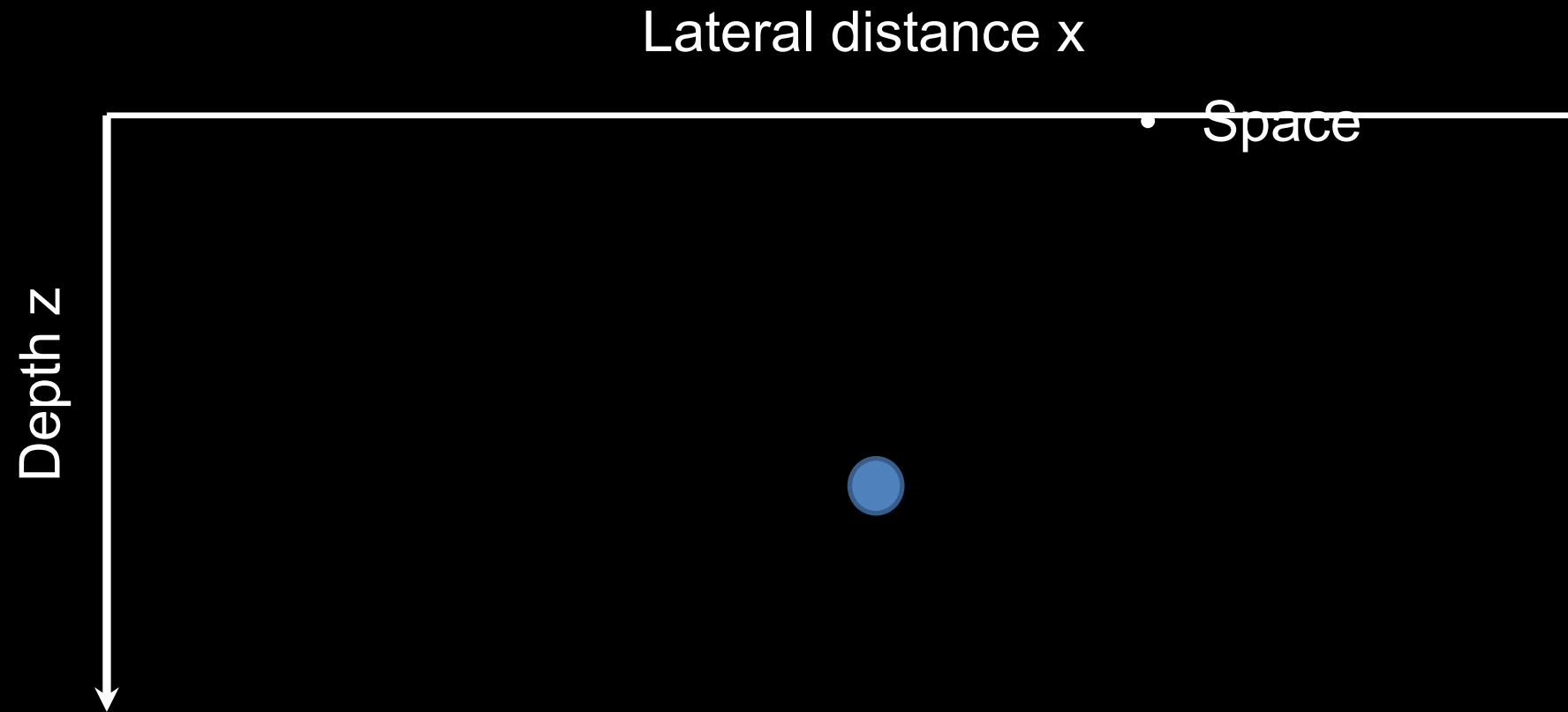
[academic use]



- Migration: What is it and why do we need it in reflection data (seismics or ground-penetrating radar) ?

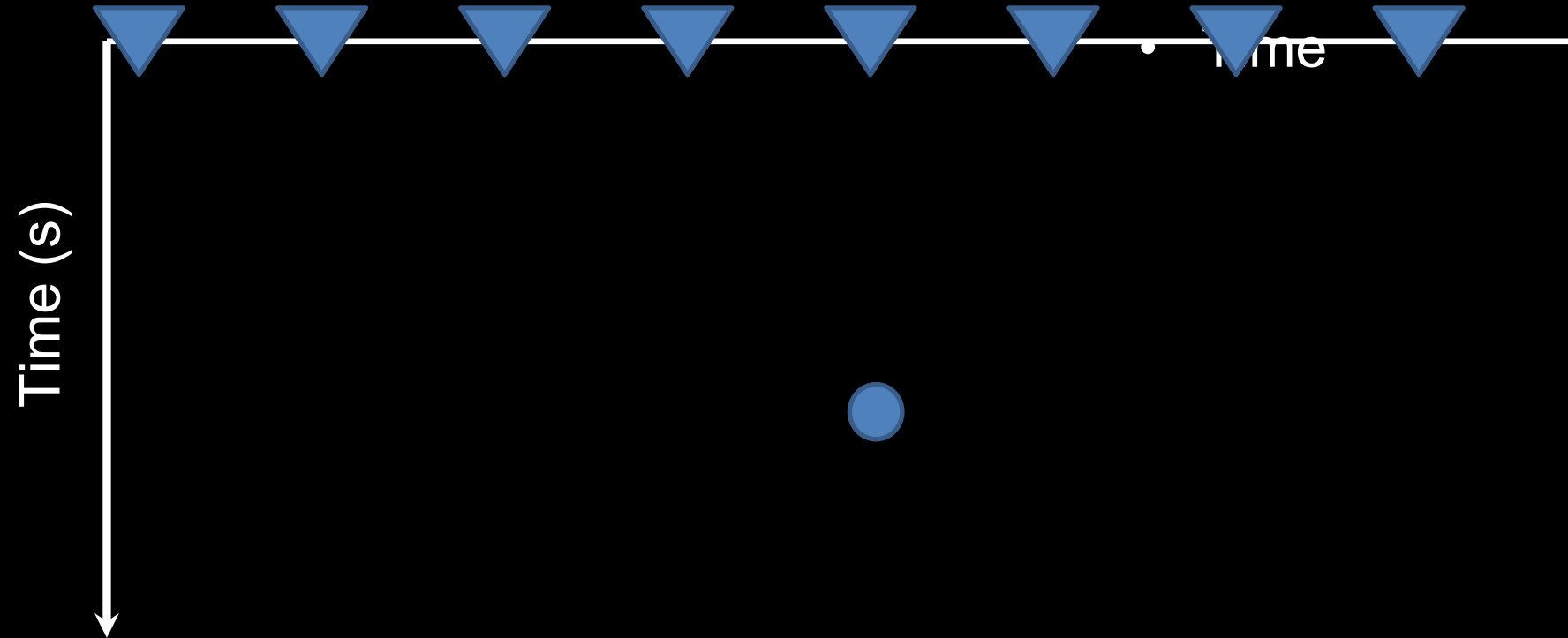


Kirchhoff Migration



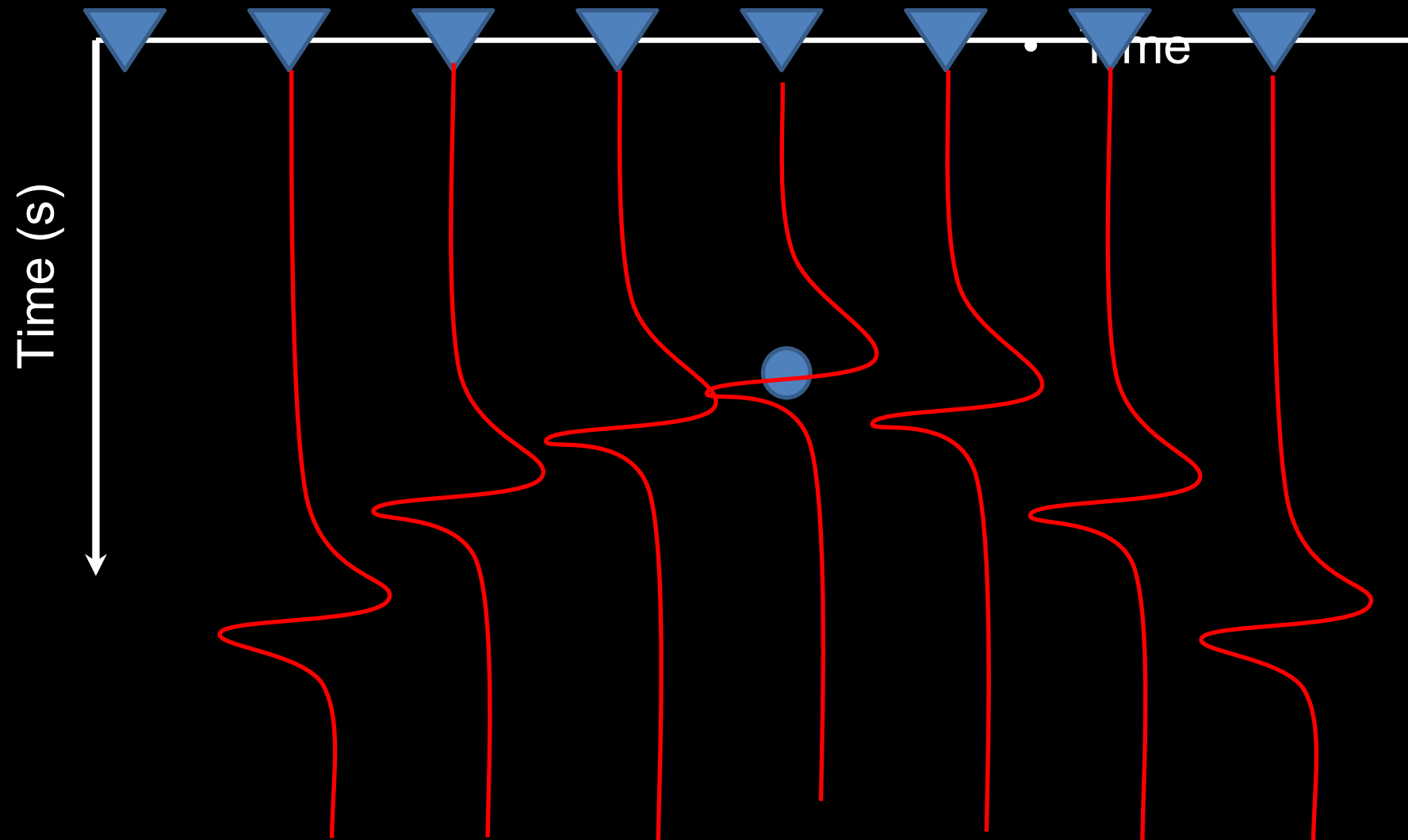


Lateral distance x Kirchhoff Migration Yilmaz 1987



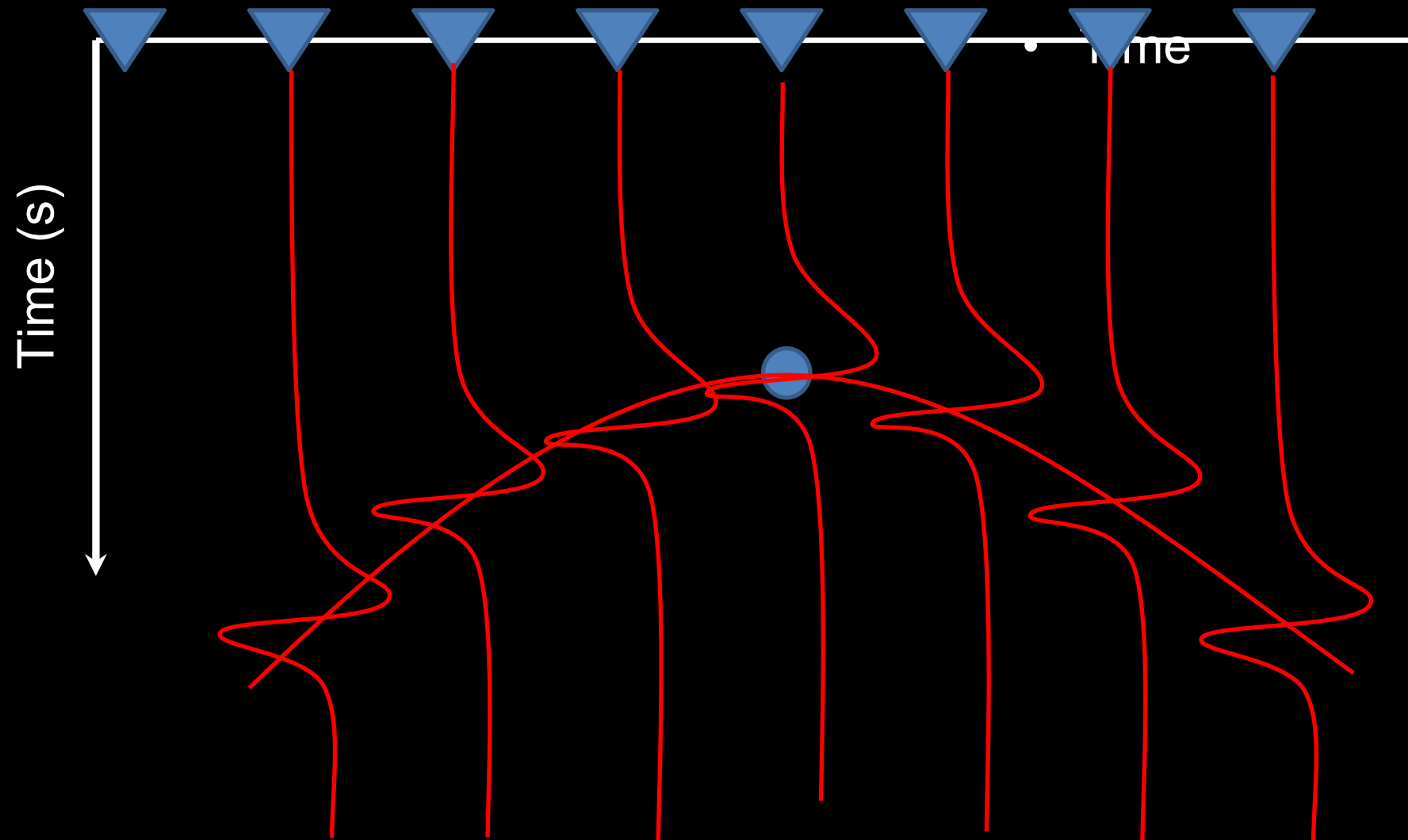


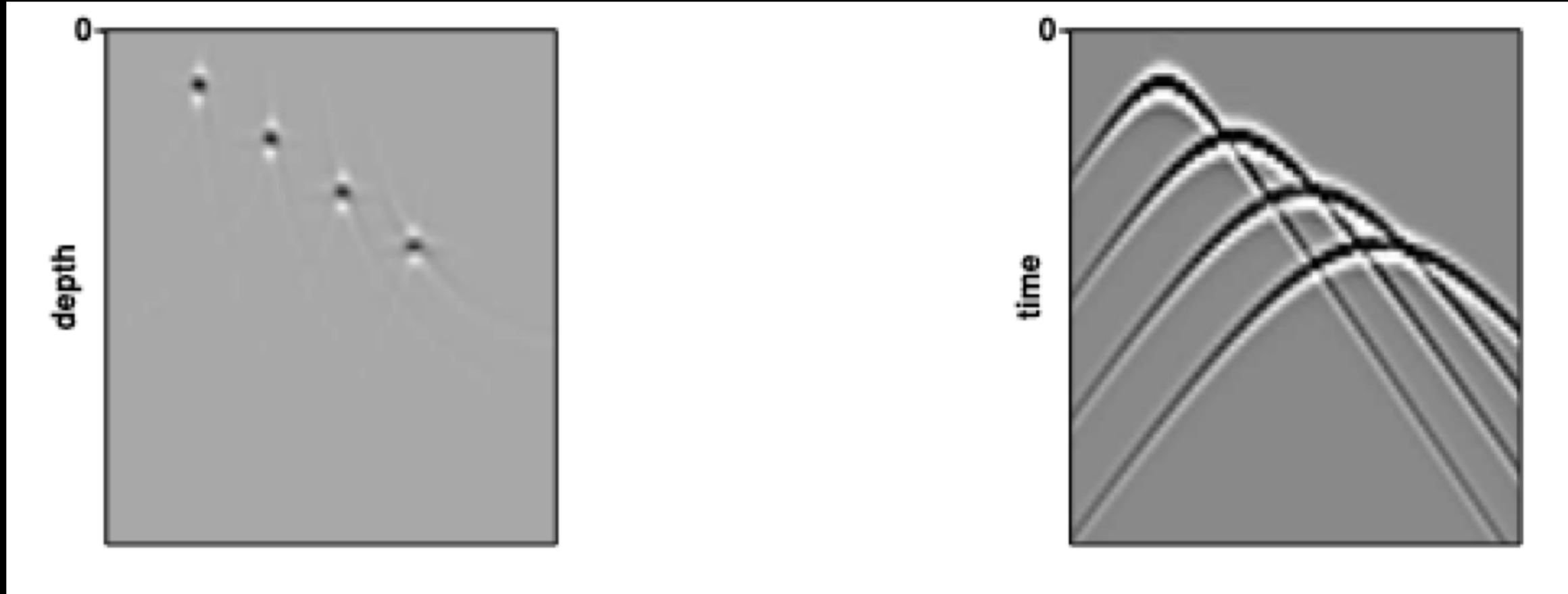
Lateral distance x Kirchhoff Migration Yilmaz 1987



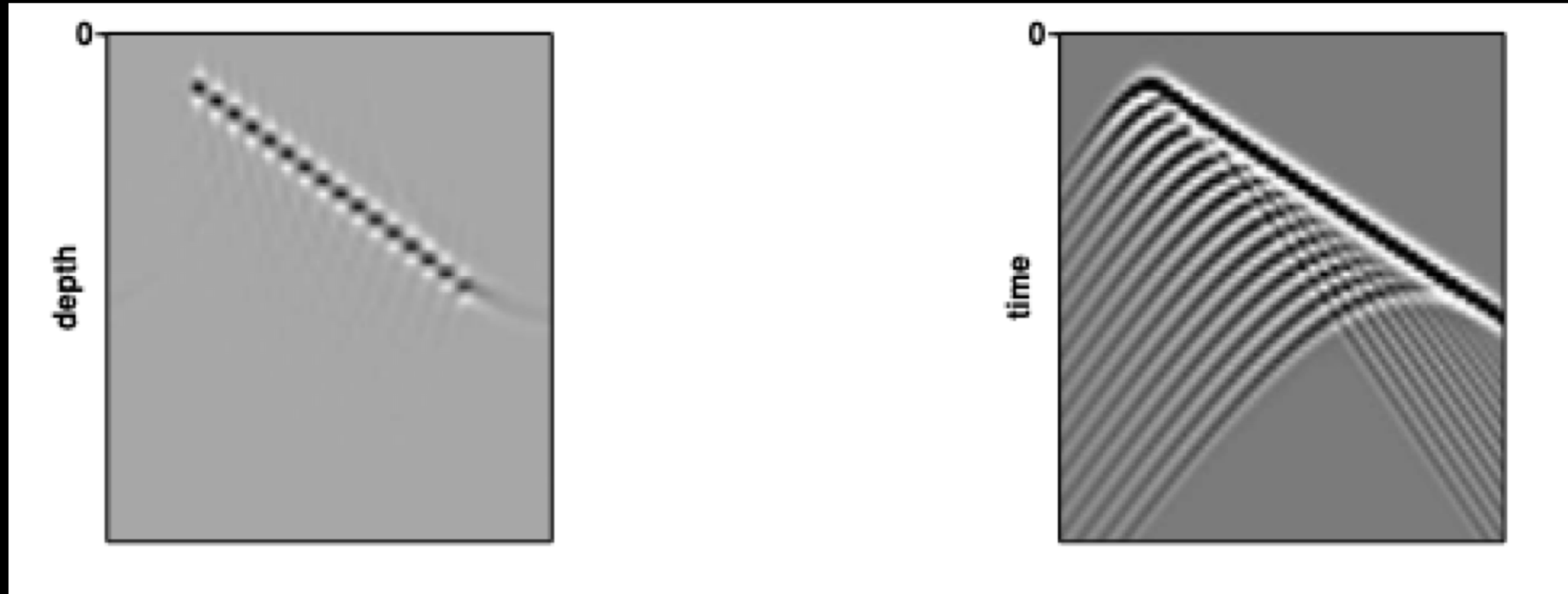


Lateral distance x Kirchhoff Migration Yilmaz 1987

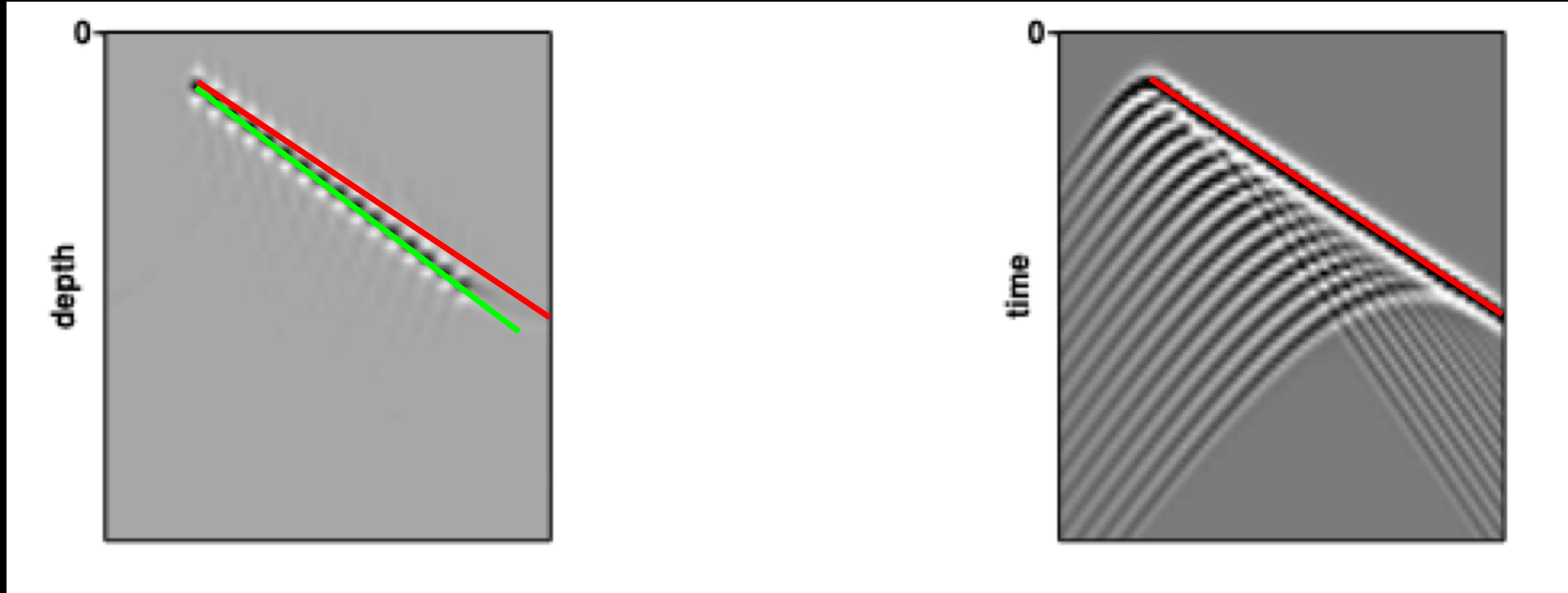




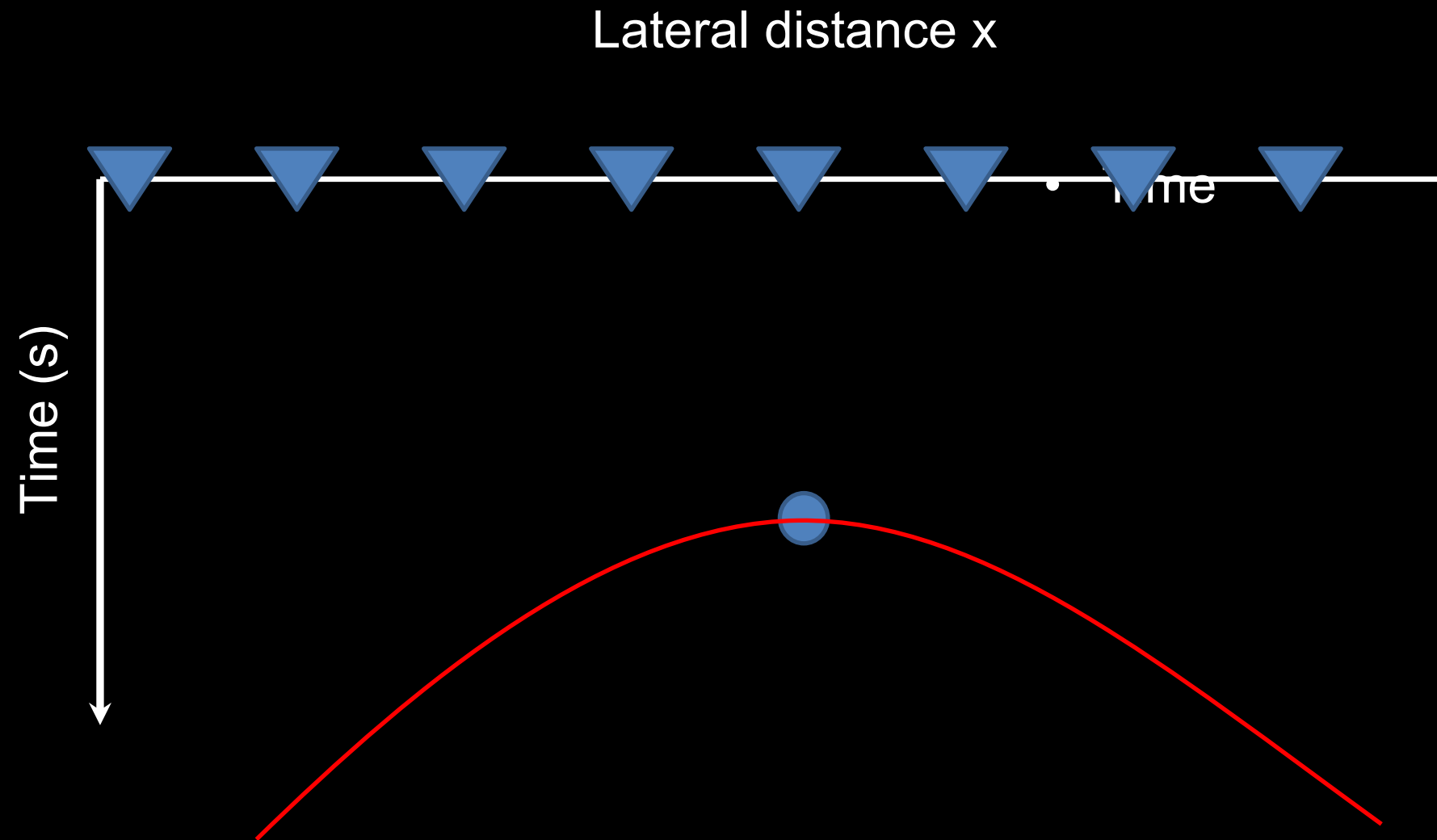
- Real World (left) vs. seismic/radar world (right)

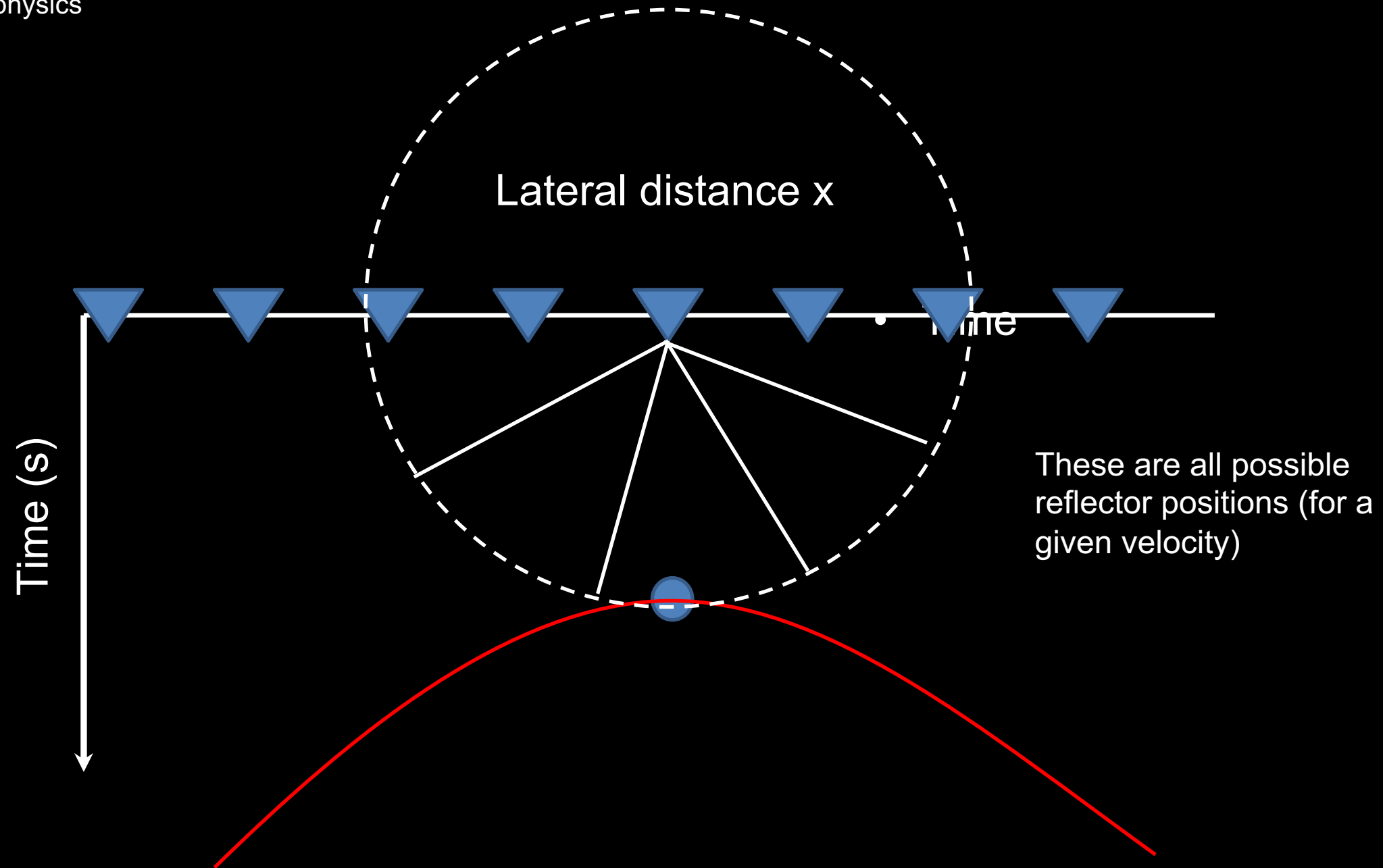


- Real World (left) vs. seismic/radar world (right)



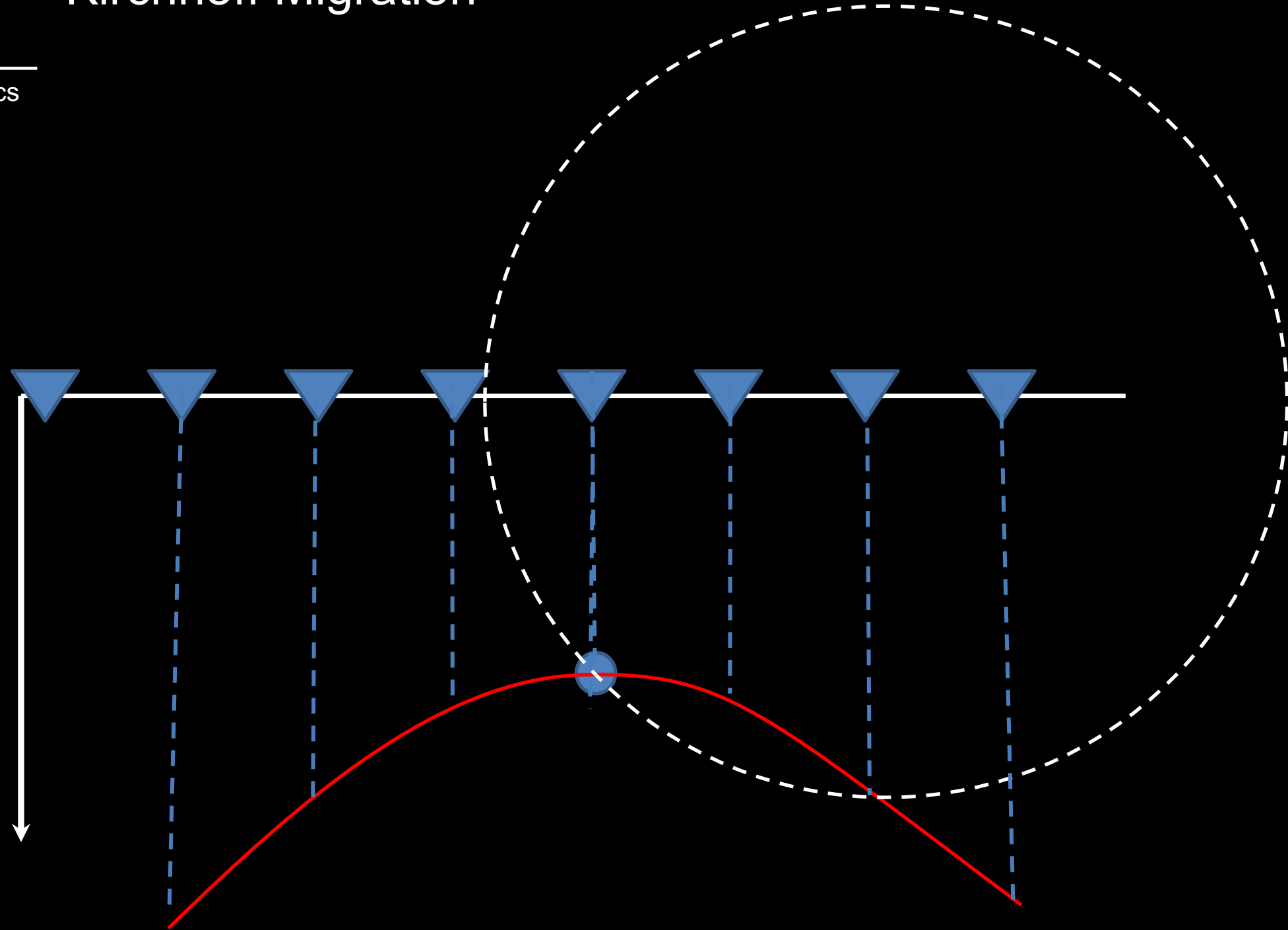
Slopes in common offset sections appear gentler than the slope of the corresponding reflection interface.

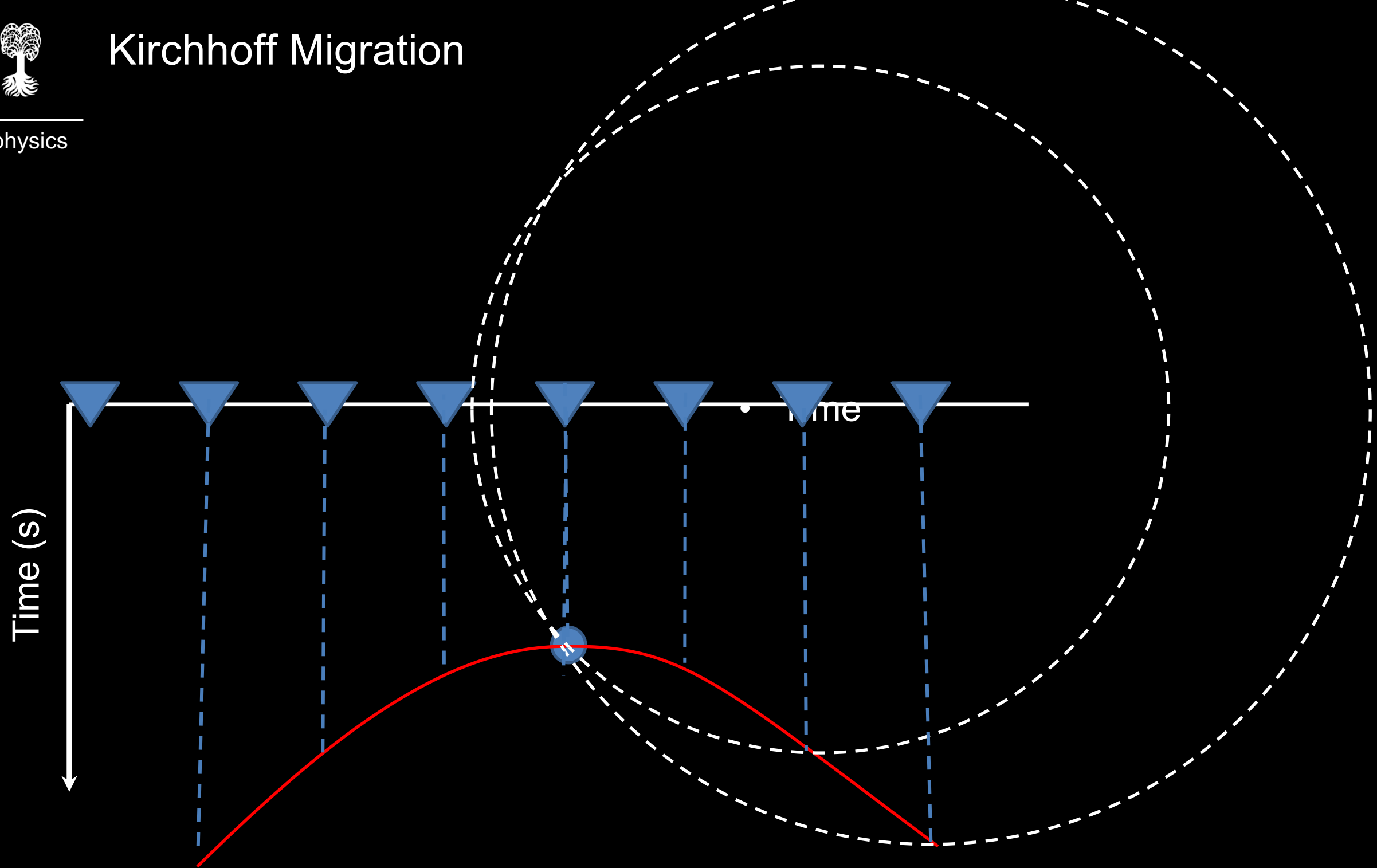






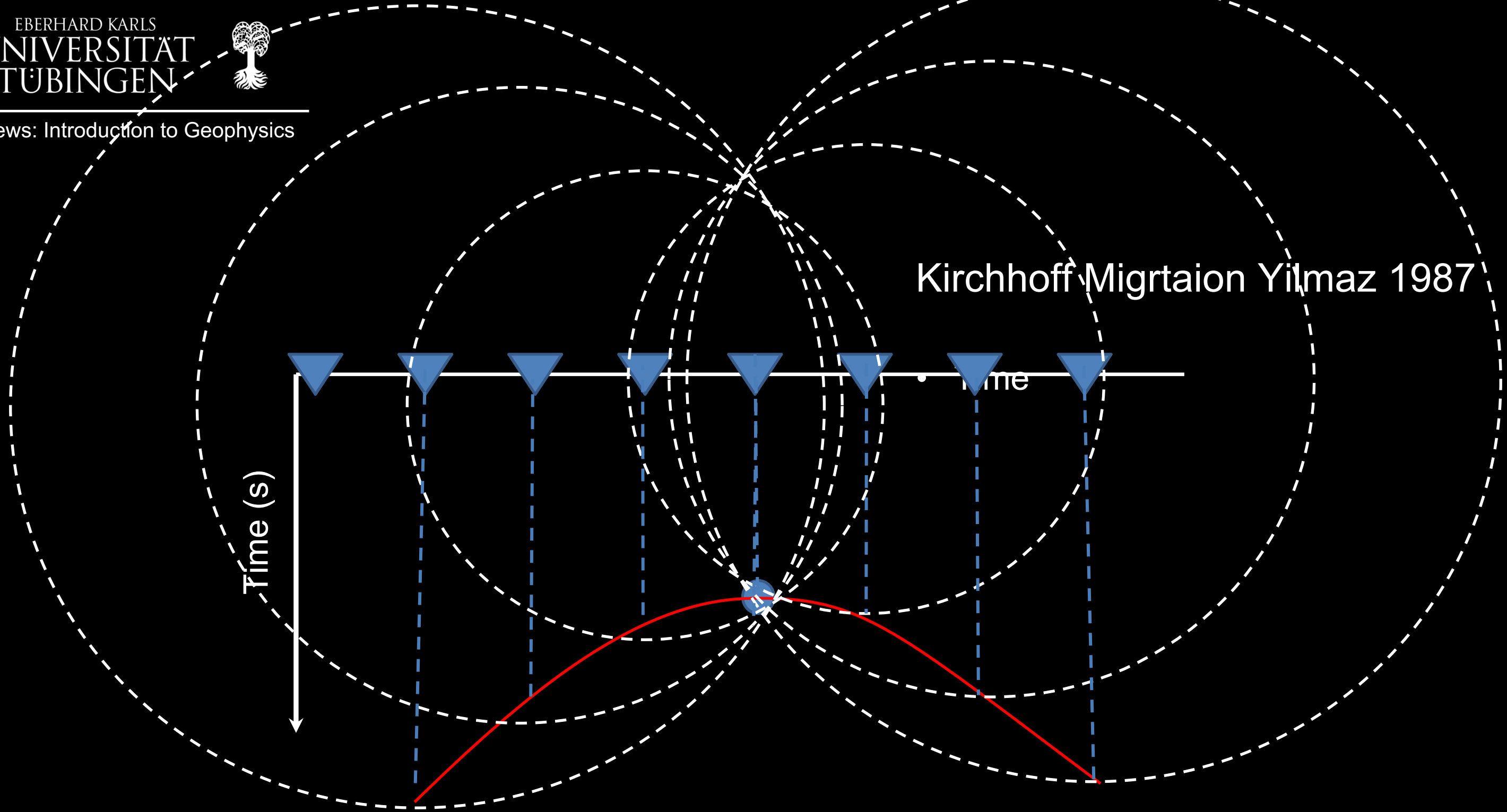
Time (s)





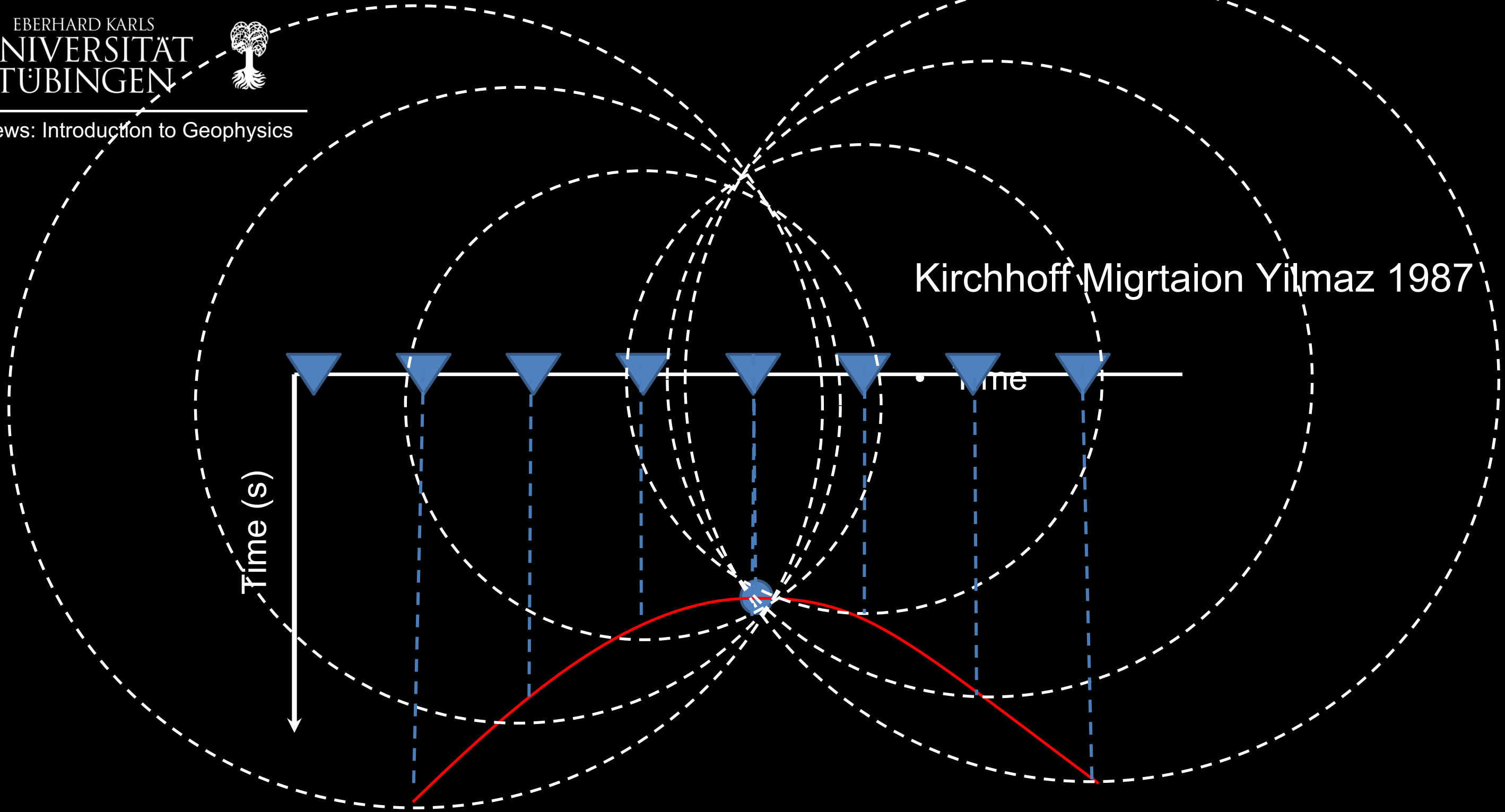


Kirchhoff Migration Yilmaz 1987





Kirchhoff Migration Yilmaz 1987





Kirchoff migration is a simple geometric way to migrate common-offset data for known velocities.

It focuses diffraction hyperbolas into points and also results in correct layer dips.

