

Logistics

- paulina.seibel@gmx.de
- Sign up today @ 18:00 - own report
- 16.06 Geophysics BBQ @ Lamswiesen (bei der Brücke B27)
- 14.05 Pint of Science Cafe Haag → evening 18:30

VL7

Electrical properties of the subsurface

Seismics: We deal with elastic properties (Young's modulus....) which are emcaps in seismic wave velocities.

Electrics: Explores the subsurface in terms of electr. conductivity/resistivity.

For now: We only focus on time-independent processes ^{direct current} CDC ^{fast phase shifts} current not AC

Basic Relationships

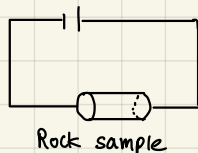
Ohm's law in an electrical circuit:

$$U = R \cdot I$$

U: Potential difference [V]

I: Current [A]

R: Resistance [Ω] \leftrightarrow Conductance $\sigma = \frac{1}{R}$ [S]

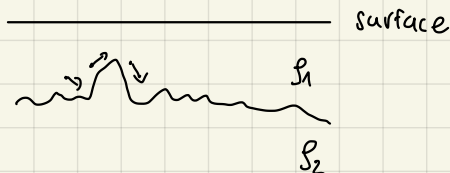


Resistance depends on geometry (A, l)

resistivity is a material property only $\frac{l}{A} \rho = R \leftrightarrow \rho = \frac{A}{l} R$ [Ωm]

Spezifischer Widerstand

For Geo/Env. application it's required to expand Ohms law for a continuous (spatially extended) medium:



$$\vec{j} = \sigma \vec{E} \quad \text{Vectorized Ohms law}$$

$\vec{j}(x, y, z)$: current density

$\sigma = \frac{1}{\rho}$: conductivity

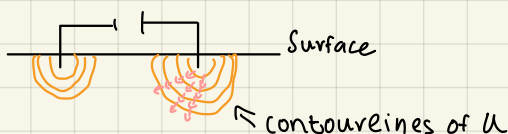
$\vec{E}(x, y, z)$: Electric field [$\frac{V}{m}$]

With $\vec{E} = -\nabla U$

$U = U(x, y, z)$ "scalar field"

"numbers"

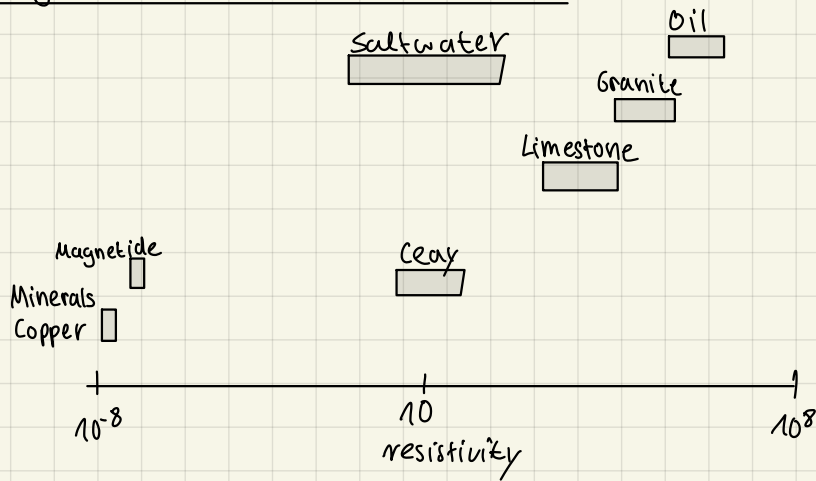
"Vector field" → number + direction



$$\vec{E} = -\nabla U \rightarrow \vec{E} \perp \text{Contour lines} \\ \rightarrow \vec{j} \parallel \vec{E}$$

σ can be a tensor → conductivity depends on direction

Range of resistivities of Earth mat.



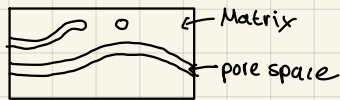
Large, large, large range !

Factors influencing conduction

here: ionic conduction

1. Conductivity of pore fluid: number of dissolved ions (Na^+ , Cl^- , ...)

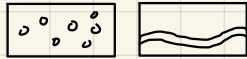
2. Porosity of rocks



more pore space \rightarrow better conductivity

3. Saturation: is the pore filled?

4. Connectedness



Archie's Law (most cited paper)

Formation factor:

$$F = \frac{\rho_0}{\rho_f} \quad \left\{ \begin{array}{l} \leftarrow \text{fully saturated rock} \\ \leftarrow \text{resistivity pore fluid} \end{array} \right.$$

Empirically:

$$F = \frac{\rho_0}{\rho_f} = \phi^{-m} \quad \left\{ \begin{array}{l} \phi: \text{porosity} \\ m: \text{most rocks} > 1 \end{array} \right\} \text{no units}$$

$\rightarrow m$ is "indicative" of degree of cementation.

\rightarrow low values of m indicate good connectedness

\rightarrow All of this is only valid for the fully saturated case. All pore space filled with brine.

$$\rho_0 = \rho_f \phi^{-m}$$

What happens if pore space is "invaded" by a different component such as oil?

If water/brine solution is displaced

$$I = \frac{\rho_f}{\rho_0} \quad \leftarrow \text{partially saturated}$$

resistivity index

$$I = S_w^{-n} \quad \left\{ \begin{array}{l} n: \text{empirical constant} \\ S_w: \text{fractional water saturation} \end{array} \right.$$

• change saturation a bit
 \rightarrow resistivity changes a lot

Combined Law:

$$\rho_r = \rho_f \phi^{-m} S_w^{-n}$$

Determine exponents in lab for a given host rock

1. We understood that resistivity is great way to characterize the subsurface because of its large range and dependency to pore fluids.

2. However, to measure any ρ/σ we need electrical currents

active
Battery
geoelectrics / resistivity
mapping

Passive
natural currents
Self-potential
(easiest & inexpensive geophysical technique)